PAPERS



# Horizontal miscible displacements through porous media: the interplay between viscous fingering and gravity segregation

Japinder S. Nijjer<sup>1</sup>,<sup>†</sup>, Duncan R. Hewitt<sup>2</sup> and Jerome A. Neufeld<sup>3,4,5</sup>

<sup>1</sup>Molecular Cellular and Developmental Biology, Yale University, 260 Whitney Avenue, New Haven 06511, USA

<sup>2</sup>Department of Mathematics, University College London, 25 Gordon Street, London WC1H 0AY, UK <sup>3</sup>Department of Applied Mathematics and Theoretical Physics, University of Cambridge, Wilberforce

Road, Cambridge CB3 0WA, UK

<sup>4</sup>Department of Earth Sciences, University of Cambridge, Downing Street, Cambridge CB2 3EQ, UK
 <sup>5</sup>BP Institute, University of Cambridge, Bullard Laboratories, Madingley Road, Cambridge CB3 0EZ, UK

(Received 7 October 2020; revised 7 November 2021; accepted 26 November 2021)

We consider miscible displacements in two-dimensional homogeneous porous media where the displacing fluid is less viscous and has a different density than the displaced fluid. We find that the dynamics evolve through nine possible regimes depending on the viscosity ratio, strength of density variations and the strength of the background flow, as characterized by the Péclet number. At early times the interface is dominated by longitudinal diffusion before undergoing a transition to a slumping regime where vertical flow is important. At intermediate times, vertical flow and diffusion can be neglected and there are three different limiting solutions: a fingering limit; an injection-driven gravity-current limit; and a density-driven gravity-current limit. Finally at late times, transverse diffusion becomes important and there is a transition from an apparent shutdown regime to a viscously enhanced Taylor-slumping regime. In each of the regimes, the dominant scalings are identified and reduced-order models for the evolution of the concentration field are developed. Lastly, three case studies are considered to illustrate the dominant physical balances in the geophysically relevant setting of geological  $CO_2$ storage.

Key words: mixing and dispersion, porous media, fingering instability

†Email address for correspondence: japinder.nijjer@yale.edu

© The Author(s), 2022. Published by Cambridge University Press. This is an Open Access article, distributed under the terms of the Creative Commons Attribution licence (https://creativecommons.org/licenses/by/4.0/), which permits unrestricted re-use, distribution, and reproduction in any medium, provided the original work is properly cited.

## 1. Introduction

Gravity currents in porous media have been studied in a wide array of contexts including the geological storage of carbon dioxide (Bickle et al. 2007; Boait et al. 2012), geothermal power generation (Woods 1999), contaminant migration (Simmons, Fenstemaker & Sharp 2001) and coastal aquifer dynamics (Fleury, Bakalowicz & de Marsily 2007). When a denser (lighter) fluid displaces a lighter (denser) fluid saturating a porous medium, the density difference causes the displacing fluid to preferentially tilt and flow along the bottom (top) boundary. Huppert & Woods (1995) studied this problem assuming the fluids had constant viscosity, the flow was purely horizontal (vertical-flow equilibrium) and there was no mixing between the fluids (sharp-interface limit). They found that the interface between the two fluids tilted, leading to self-similar spreading of the fluids. More recently, the effects of diffusion and vertical flow were considered by Szulczewski & Juanes (2013). When the two fluids were fully miscible and vertical flow was accounted for, they found that a series of different regimes arose depending on the dominant physical balances. At intermediate times they found that diffusion and vertical flow could be neglected and the flow reproduced the dynamics outlined by Huppert & Woods (1995); however, at early and late times, diffusion played a dominant role in setting the spreading rate.

In addition to having different densities, the two fluids may also have different viscosities, a complication neglected by Huppert & Woods (1995) and Szulczewski & Juanes (2013). Hesse, Tchelepi & Cantwell (2007), Pegler, Huppert & Neufeld (2014) and Zheng et al. (2015) extended the work of Huppert & Woods (1995) to study the effect of differing viscosities and densities on the evolution of the displacement front in the limit of sharp interfaces and vertical-flow equilibrium. They found that when the height of the current was comparable to the height of the medium, the viscosity ratio played a dominant role in setting the spreading rate of the gravity current, leading to significantly enhanced spreading if the injected fluid was less viscous than the ambient fluid, or reduced spreading if the injected fluid was more viscous than the ambient fluid. However, similar to Huppert & Woods (1995), these works also omit the effects of diffusion and vertical flow. Furthermore, these studies also assumed that the interface between the two fluids was hydrodynamically stable. However, when a less-viscous fluid displaces a more-viscous fluid, viscous fingering can develop and lead to the interpenetration of the two fluids and a highly non-trivial interface (Homsy 1987). Thus far, the effect of viscous fingering on spreading in depth-integrated gravity current models has not been well understood.

In a similar vein, a number of authors have looked at the effect of gravity on the viscous-fingering instability: Rogerson & Meiburg (1993*b*) studied the onset of the viscous-fingering instability with a gravitationally driven shear parallel to the interface using linear stability analysis; Rogerson & Meiburg (1993*a*), Tchelepi *et al.* (2004), Tchelepi & Orr (1994), Ruith & Meiburg (2000), Camhi, Meiburg & Ruith (2000) and Riaz & Meiburg (2003) investigated the nonlinear evolution of the fingering instability using numerical simulations; and Tchelepi & Orr (1994), Berg *et al.* (2010) and Suekane, Koe & Barbancho (2019) examined the nonlinear evolution of the fingering instability using laboratory experiments. While these studies highlight some of the interesting qualitative behaviour that can be observed, they do not provide a full overview of the different dynamical regimes, nor do they provide quantitative predictions for the evolution of the concentration field. Furthermore, in all of the work discussed above, the long-time asymptotic behaviour, where diffusion and mixing are important, is neglected.

The overarching aim of this work is to bridge the gap between these two different bodies of work in order to develop a quantitative understanding of the full life-cycle of miscible viscous fingering with gravitational segregation. This work is laid out as follows. In § 2,

we formulate the problem and in § 3 consider the two limiting cases of uniform viscosity and uniform density. In § 4 the effects of both density and viscosity differences are examined and the flow phenomenology and dominant dynamical regimes are identified; in §§ 4.1-4.3, these regimes are examined in more detail and reduced-order models for the evolution of the concentration field are derived. Finally, in § 5, the results are summarized and the implications of the results for carbon capture and geological storage are briefly discussed.

## 2. Problem formulation

We consider a semi-infinite 2-D porous layer with a uniform permeability k (figure 1). A fluid with density  $\rho_1$  and viscosity  $\mu_1$  is injected, with a volume flux Q, into the medium, which is saturated with an ambient fluid of density  $\rho_2$  and viscosity  $\mu_2$ . We assume that the fluid flow obeys Darcy's law, that the flow is incompressible, and that the concentration c of the injected fluid evolves through the action of advection and diffusion:

$$\boldsymbol{u} = -\frac{k}{\mu(c)} \left( \nabla p + \rho g \hat{\boldsymbol{y}} \right), \tag{2.1}$$

$$\nabla \cdot \boldsymbol{u} = \boldsymbol{0}, \tag{2.2}$$

$$\phi \frac{\partial c}{\partial t} + \boldsymbol{u} \cdot \boldsymbol{\nabla} c = \phi D \boldsymbol{\nabla}^2 c.$$
(2.3)

Here u = (u, v) is the Darcy velocity, p the pressure and D an isotropic and constant diffusion coefficient. Note that the injected concentration,  $C_i \neq 1$ , and the ambient concentration  $C_a \neq 0$ , can be accounted for with a simple linear rescaling. Following the convention of previous works (Tan & Homsy 1988; Camhi *et al.* 2000), we assume the viscosity varies exponentially with the concentration

$$\mu(c) = \mu_2 e^{-Rc}, \tag{2.4}$$

where  $R = \log(\mu_2/\mu_1)$  is the log-viscosity ratio. We employ the Boussinesq approximation and assume a linear relationship between the concentration and the density  $\rho$ ,

$$\rho = \rho_2 + (\rho_1 - \rho_2)c. \tag{2.5}$$

### 2.1. Non-dimensionalization

We change to a reference frame moving with the average injection velocity with transformed coordinate  $\tilde{x} = x - Qt/a$  and velocity  $\tilde{u} = u - Q/a$ . We also incorporate the hydrostatic contributions of the ambient fluid into the pressure field:  $\tilde{p} = p + \rho_2 gy$ . Non-dimensionalizing by the width of the medium *a*, velocity Q/a, advective time scale  $\phi a^2/Q$ , viscosity  $\mu_2$  and pressure  $\mu_2 Q/k$ , equation (2.1)–(2.5) become

$$-(\tilde{u}^*+1)\mu^* = \frac{\partial \tilde{p}^*}{\partial \tilde{x}^*}, \quad -v^*\mu^* = \left(\frac{\partial \tilde{p}^*}{\partial y^*} + Gc\right), \quad (2.6a,b)$$

$$\frac{\partial \tilde{u}^*}{\partial \tilde{x}^*} + \frac{\partial v^*}{\partial y^*} = 0, \qquad (2.7)$$

935 A14-3



Figure 1. A schematic of the model geometry. The porous medium is infinite in the  $\hat{x}$  direction, and finite, with width a, in the  $\hat{y}$  direction. The medium is initially saturated with a fluid of density  $\rho_2$  and viscosity  $\mu_2$  and another fully miscible fluid, with density  $\rho_1$  and viscosity  $\mu_1$  is injected at the left boundary. We assume that the injected fluid is introduced in a purely horizontal manner and at a constant two-dimensional (2-D) volumetric flow rate Q.

$$\frac{\partial c}{\partial t^*} + \tilde{u}^* \frac{\partial c}{\partial \tilde{x}^*} + v \frac{\partial c}{\partial y^*} = \frac{1}{Pe} \left( \frac{\partial^2 c}{\partial \tilde{x}^{*2}} + \frac{\partial^2 c}{\partial y^{*2}} \right), \tag{2.8}$$

$$\mu^* = \mathrm{e}^{-Rc},\tag{2.9}$$

where (\*) denotes a dimensionless quantity. For notational convenience we drop the tildes and asterisks in all subsequent expressions.

There are three important non-dimensional parameters in this problem:

$$R = \log(\mu_2/\mu_1), \quad Pe = \frac{Q}{D}, \quad G = \frac{g(\rho_1 - \rho_2)ka}{Q\mu_2}.$$
 (2.10*a*-*c*)

The log-viscosity ratio R, measures the strength of the viscosity variations. We will only consider the case R > 0, that is, when the injected fluid is less viscous than the ambient fluid. The Péclet number, Pe, measures the relative strength of advection to diffusion. In this work, we will focus predominantly on the geologically relevant limit,  $Pe \gg 1$ . Note that the Péclet number defined here is a macroscopic quantity that depends on the overall size of the porous medium. It is much larger than the pore-scale Péclet number  $Pe_p = Qb/aD$  which is calculated based on the intrinsic pore scale, b, of the porous medium.

Finally, the gravity number, G, measures the ratio of pressure gradients due to density differences and those due to injection. Alternatively, it can be interpreted as a ratio of the vertical rise/fall velocity due to gravity and the horizontal injection velocity. We only consider G > 0, since G < 0 is equivalent to G > 0 with a vertical reflection of the coordinate system  $y \rightarrow -y$ . Note that the strength of the gravitational force is sometimes defined in terms of a Rayleigh number,  $Ra = g(\rho_2 - \rho_1)ka/D\mu_2 = GPe$ .

#### 2.2. Boundary conditions

The fluid is injected with a constant horizontal flux, and so in the moving frame

$$\frac{\partial c}{\partial x} \to 0 \quad \text{as } x \to \pm \infty,$$
 (2.11)

$$\int_0^1 u \, \mathrm{d}y \to 0 \quad \text{as } x \to \pm \infty, \tag{2.12}$$

935 A14-4

https://doi.org/10.1017/jfm.2021.1067 Published online by Cambridge University Press

$$v \to 0 \quad \text{as } x \to \pm \infty.$$
 (2.13)

The last constraint is equivalent to a hydrostatic far-field pressure. There is no flux of concentration through the impermeable upper and lower boundaries of the channel, and so

$$\frac{\partial c}{\partial y} = 0 \quad \text{for } y = 0, 1, \tag{2.14}$$

$$v(x, 0, t) = v(x, 1, t) = 0.$$
 (2.15)

No-flux boundaries are critical to allow for the formation of currents that propagate along the boundaries (Rogerson & Meiburg 1993a; Ruith & Meiburg 2000). We initialize the concentration field to have a step jump,

$$c(x, t = 0) = c_0(x) = H(-x),$$
(2.16)

where H(x) is the Heaviside function.

At each time step the velocity field is found using a multigrid relaxation method (Adams 1999). The concentration field is advanced in time using a third-order Runge–Kutta scheme, the advective term  $\boldsymbol{u} \cdot \nabla c$  is discretized using a third-order upwinding scheme and the diffusive term  $\nabla^2 c/Pe$  is discretized using a sixth-order compact finite difference method.

## 2.3. Diagnostic quantities

In order to investigate how the large-scale flow evolves, we focus on quantifying and predicting the evolution of transversely averaged concentration  $\bar{c}(x, t) = \int_0^1 c \, dy$ . Correspondingly, the concentration deviations are defined as  $c'(x, y, t) = c(x, y, t) - \bar{c}(x, t)$ . Integrating (2.8) over the depth of the porous strip and noting that the flow is incompressible and there is no flux through the top and bottom boundaries yields

$$\frac{\partial \bar{c}}{\partial t} + \frac{\partial \overline{uc'}}{\partial x} = \frac{1}{Pe} \frac{\partial^2 \bar{c}}{\partial x^2}.$$
(2.17)

Subtracting (2.17) from (2.8) gives the evolution equation for the deviations

$$\frac{\partial c'}{\partial t} + \frac{\partial (uc')}{\partial x} + \frac{\partial (u\bar{c})}{\partial x} - \frac{\partial (\overline{uc'})}{\partial x} + \frac{\partial (vc)}{\partial y} = \frac{1}{Pe} \left( \frac{\partial^2 c'}{\partial y^2} + \frac{\partial^2 c'}{\partial x^2} \right).$$
(2.18)

We also examine two macroscopic quantities in time: the spreading length h(t) (also referred to as the 'mixing' length) and the Nusselt number Nu. We define the spreading length, which measures the length over which the two fluids have spread, as the second spatial moment of the concentration about the initial condition  $c_0(x)$  as defined in (2.16),

$$h(t) = \sqrt{\frac{\int_{-\infty}^{\infty} x^2 (\bar{c} - c_0)^2 \,\mathrm{d}x}{\int_{-\infty}^{\infty} (\bar{c} - c_0)^2 \,\mathrm{d}x}}$$
(2.19)

(cf. Nijjer, Hewitt & Neufeld 2019). To quantify the rate at which the ambient and injected fluids spread, we define the advective mass exchange through the midplane (travelling with

the mean injection velocity),

$$Nu = \int_0^1 (uc)|_{x=0} \,\mathrm{d}y, \tag{2.20}$$

which represents the rate at which the concentration is advectively exchanged between the two fluids. When Nu = 0, there is no advective exchange: the two fluids simply advect to the right at the injection speed and the interface between the fluids only evolves by diffusion. Larger values of Nu indicate more advective exchange across the moving interface. Note that the Nusselt number is often used to quantify the total (advective plus diffusive) transport but here we use it to quantify just the advective transport.

Although not used as a diagnostic quantity, it is useful to combine (2.6*a*,*b*), (2.7) and (2.9) and write the velocity in terms of a stream function  $(u, v) = (\partial \Psi / \partial y, -\partial \Psi / \partial x)$ , as follows:

$$\frac{\partial^2 \Psi}{\partial x^2} + \frac{\partial^2 \Psi}{\partial y^2} - R \frac{\partial \Psi}{\partial x} \frac{\partial c}{\partial x} - R \frac{\partial \Psi}{\partial y} \frac{\partial c}{\partial y} = R \frac{\partial c}{\partial y} + G e^{Rc} \frac{\partial c}{\partial x}.$$
 (2.21)

## 3. Limiting cases

In this section, we consider two limiting cases: the uniform-viscosity case (log-viscosity ratio R = 0) where the fluids differ only in density; and the uniform-density case (gravity number G = 0) where the fluids differ only in viscosity.

## 3.1. Uniform viscosity, R = 0

When R = 0 the problem reduces to the uniform-viscosity gravity-current problem. This situation was studied by Szulczewski & Juanes (2013) and we give a brief overview of the dynamics here in order to frame our subsequent more general results. Snapshots of the concentration field from a representative simulation are given in figure 2(a-e). In general, the density difference between the two fluids causes the interface to tilt, with the denser injected fluid travelling along the bottom boundary, which aids in spreading the two fluids. To quantify this spreading, we plot the evolution of the spreading length h in figure 2(f) along with the corresponding theoretical predictions. In general, the spreading length grows in time through five different regimes, each with different power-law growth rates.

First, the concentration field is vertically homogeneous, and longitudinal diffusion dominates, leading to  $h \sim (t/Pe)^{1/2}$  (figure 2a, f). Second, once longitudinal diffusion becomes unimportant relative to advection, occurring at a time  $t \sim O(1/G^2 P e)$ , the interface slumps due to gravity. Vertical flow in this regime is important since the vertical length scale of the spreading zone is initially larger than the horizontal scale. This leads to so-called S-shaped slumping and linear growth of the spreading length,  $h \sim Gt$ (figure 2b, f). Third, once the interface has become long and thin, that is the horizontal extent of the spreading zone becomes much larger than the vertical extent, occurring at a time  $t \sim O(1/G)$ , vertical flow becomes unimportant so that the flow is predominantly horizontal. The interface continues to slump due to gravity and takes on a characteristic straight-line profile (figure 2c). The spreading length grows sublinearly,  $h \sim (Gt)^{1/2}$ , since the hydrostatic pressure gradient, which drives the flow, diminishes as the interface slumps (figure 2f). Fourth, once the interface has become even longer and thinner, occurring at a time O(Pe), transverse diffusion becomes important. A balance between horizontal advection and transverse diffusion results in net horizontal transport analogous to Taylor dispersion (Taylor 1953). However, because the horizontal velocity is proportional to the



Figure 2. Evolution of the concentration field for R = 0 and (G, Pe) = (3, 100). (a-e) Plots of the concentration field versus x/h(t) and y at  $(a) t = 1 \times 10^{-4}$ , (b) t = 0.036, (c) t = 0.83, (d) t = 57 (e) t = 8000. (f) Evolution of the spreading length, h, as a function of time, t. The dots correspond to the snapshots in panels (a-e) and the dashed lines correspond to the theoretical predictions from Szulczewski & Juanes (2013).

horizontal gradient in the concentration, the transport is subdiffusive,  $h \sim (G^2 Pet)^{1/4}$  (figure 2d, f). Fifth, once the shear-enhanced dispersivity becomes small compared with molecular diffusion, occurring at a time  $t \sim O(G^2 Pe^3)$ , the interface grows diffusively again,  $h \sim (t/Pe)^{1/2}$  (figure 2e, f). Note that in each case, the transition times are found by determining the time at which the spreading lengths overlap.

## 3.2. Uniform density, G = 0

When G = 0, the problem reduces to the miscible viscous-fingering problem studied by Nijjer *et al.* (2018). To review the results of the earlier work, at early times the interface grows diffusively,  $h \sim (t/Pe)^{1/2}$ , while the instability grows exponentially (figure 3a, f). At intermediate times, the instability saturates and fingers elongate and interact nonlinearly leading to coarsening and advective growth of the spreading length,  $h \sim Rt$  (figure 3b, f). At late times, a single dominant finger is left in the centre of the domain with counter-propagating fingers along the top and bottom boundaries (figure 3c) that eventually slow, leaving a well-mixed interior with constant  $h \rightarrow RPe$  (figure 3d, f). Over very long times, in the absence of any advection, the fluid interface evolves purely through longitudinal diffusion and so the spreading length grows like  $h \sim (t/Pe)^{1/2}$ .

One subtle difference between the problem studied in Nijjer *et al.* (2018) and the one studied here is that no-penetration conditions are imposed here along the top and bottom boundaries instead of periodic boundary conditions. We find that up until the late-time regime, this difference in boundary conditions has little effect on the dynamics. However, at late times as the single-finger exchange-flow decays, a pair of wider counter-propagating fingers manifest themselves along the boundaries (figure 3e, f). This is because, in contrast to the case with periodic boundaries, a half-wavelength mode is now permissible (Abdul Hamid & Muggeridge 2020). Since this mode is wider, it decays more slowly and is still unstable once the central propagating finger decays away. This mode decays four times more slowly and spreads four times farther but also eventually decays away.



Figure 3. Evolution of the concentration field for G = 0 and (R, Pe) = (2.5, 500). (a-e) Plots of the concentration field versus x/h(t) and y at  $(a) t = 2 \times 10^{-2}$ , (b) t = 2.6, (c) t = 15, (d) t = 170 (e) t = 525. (f)Evolution of the spreading length, h, as a function of time, t. The dots correspond to snapshots in panels (a-e) and the dashed lines correspond to the theoretical predictions from Nijjer, Hewitt & Neufeld (2018).

The temporal scalings for the early- to intermediate-time transition and intermediateto late-time transitions are as before, occurring at times  $t \sim O(1/R^2Pe)$  and  $t \sim O(Pe)$ , respectively, again calculated by determining the overlap of the spreading lengths. There is an additional time transition to the half-wavelength mode which occurs at a time  $t \sim O(Pe)$ as well, but with a larger prefactor.

## 3.3. Comparison of the two limiting cases

There are a number of similarities between the two limiting cases discussed. In both cases, the early-time dynamics are dominated by longitudinal diffusion. At intermediate times diffusion becomes unimportant and spreading is dominated by advection. Initially, vertical flow is important but, as the interface is stretched longitudinally, the flow becomes predominantly horizontal. At late times, there is a balance between horizontal advection and transverse diffusion leading to a slowdown in the flow.

Despite these similarities, the manner in which the systems evolve are very different. In the uniform-viscosity case, there are no hydrodynamic instabilities and the dynamics are insensitive to small changes in the initial conditions, while in the uniform-density case, the interface fingers chaotically and the exact dynamics are highly sensitive to the initial conditions. At intermediate times, in the uniform-viscosity case, the spreading length first grows like  $h \sim t$  then like  $h \sim t^{1/2}$  while in the uniform-density case the spreading length grows like  $h \sim t$ . At late times, the spreading length grows like  $h \sim t^{1/4}$  in the uniform-viscosity case, the dynamics are decoupled and independent of the injection flux, whereas in the uniform-density case, injection is critical to the formation of fingers. The aim of the remainder of this work is to outline how the aforementioned similarities and differences evolve as both the viscosity and density are varied away from the limiting cases.



Figure 4. Colourmaps of the concentration field for (R, Pe) = (2, 500) and (a, c, e, g) G = 0.025, (b, d, f, h)G = 2. The snapshots are taken at times (a, b) t = 0.05, (c, d) t = 2, (e, f) t = 30 and (g, h) t = 500. Evolution of (i) the spreading length, h and (j) the Nusselt number, Nu, for the same parameters as in (a-h).

## 4. Overall dynamics

Consider the case where both the density and viscosity vary ( $G \neq 0, R \neq 0$ ). Depending on the choice of parameters, a range of different behaviours are possible. In figure 4, we show a series of snapshots in time of the concentration field for a large and a small value of G, each showing a range of different dynamics. In both cases the interface is stretched longitudinally and eventually becomes transversely well mixed, but the manner in which the flows reach this final state depend on G.

At very early times, in both cases, molecular diffusion dominates the dynamics. The interface then slumps due to gravity, forming a pair of tongues along the top and bottom boundaries, with the effect being more pronounced for larger G (figure 4a,b). The slumping is asymmetric, due to the viscosity difference, as the forward-propagating tongue propagates much faster than the backward-propagating tongue (figure 4b). At intermediate times, depending on the relative magnitudes of G and R (and Pe), the interface may finger (figure 4c) or not finger (figure 4d). When the interface fingers, the fluid spreads more slowly as compared with the non-fingered case (figure 4i,j). Eventually, the fingered interface coarsens until a pair of counter-propagating currents remain which resemble the non-fingered case (figure 4e, f). At late times, in both cases, the interface becomes vertically homogenized and the growth of the spreading length slows and tends to the same value independent of G (figure 4g,h). This is analogous to the shutdown

of the viscous fingering instability in the G = 0 limit. Eventually, horizontal advective transport becomes dominated by Taylor dispersion, analogous to the R = 0 limit, driven by horizontal gradients in *c*. This too becomes negligible over very long times and molecular diffusion dominates again.

Figure 4(*i*) shows the evolution of the spreading length for large *G*, small *G* and zero *G*. When *G* is large, the spreading length initially grows like *t*, in a manner analogous to the slumping regime in the uniform-viscosity case. The spreading length then grows with a scaling exponent less than unity before tending to a constant spreading length. When *G* is small, the dynamics are similar to the uniform-density case G = 0: the spreading length initially grows diffusively, then advectively, before tending to a constant. However, the nonlinear regime is reached earlier, the prefactor in the fingering regime is larger, and the fingers coarsen directly to the half-wavelength mode. In general, increasing *G* leads initially to faster growth, but all three examples tend to nearly the same constant spreading length  $h \sim t^{1/4}$  (not shown), with a prefactor that increases with *G* and over even longer times the spreading length grows due to longitudinal diffusion and  $h \sim t^{1/2}$ , independent of *G*.

Figure 4(*j*) shows the evolution of the advective flux through the midplane, *Nu*. When *G* is large, the advective flux starts at a maximum and initially decays slowly. After approximately t = 30, the decay rate increases, characteristic of exponential decay of the flux, before tending to a constant. When *G* is small, the advective flux initially exhibits power-law growth, then exponential growth, before saturating and fluctuating about a constant value. Eventually the flux coincides with the large *G* case and decays in the same manner, before tending to a smaller constant. The constant flux corresponds to the shear-enhanced dispersive flux driven by density differences between the two fluids. Over very long times (not shown), the advective flux decays to zero as the interface becomes more stretched. For comparison we show the uniform-density case G = 0, which grows exponentially, fluctuates about a constant and exponentially decays, before growing again as the single finger state becomes unstable as discussed in § 3.2.

In the following sections we discuss the different regimes that are possible in more detail.

#### 4.1. Early-time slumping regime

Initially, the streamwise concentration gradient between the fluids is large and the concentration is transversely homogeneous. In this case, diffusion across the interface dominates and the concentration evolves diffusively:

$$c = \bar{c} = \frac{1}{2} + \frac{1}{2} \operatorname{erf}\left(-\frac{x}{\sqrt{4t/Pe}}\right).$$
 (4.1)

Once advection begins to outpace diffusion, the interface slumps along the boundaries. This leads to localized regions of fast flow but little motion away from the boundaries. When *G* is small, small buoyancy-driven fingers, comparable to the fingering instability, grow along the top and bottom boundaries while fingers along the rest of the interface grow more slowly (figure 5*a*). This preference for fingering along the boundaries occurs because the difference in density leads to slumping which preferentially perturbs the instability along the boundaries. In this case the density difference perturbs the interface but the growth of the fingers is still dominated by viscous effects. We therefore expect that the spreading length grows in a manner analogous to the viscous fingering instability, and so  $h \sim Rt$  and a transition time from the diffusive regime at  $t \sim (R^2 P e)^{-1}$  (cf.

Interplay of viscous fingering and gravitational segregation



Figure 5. Colourmaps of the concentration field in the slumping regime for (R, Pe) = (1, 4000) and (G, t) = (a) (0.02, 0.6) and (b) (2, 0.07).



Figure 6. Plots of h(t) for Pe = 100 and different *G* and *R* as labelled. The raw data is plotted in (*a*). In (*b*) the spreading length is rescaled by the predicted scalings for the small-*G* slumping limit. The dotted lines denote simulations with G = 0 and  $R = \{0.5, 1, 2\}$ . In (*c*) the spreading length is rescaled by the predicted scalings for the large-*G* slumping limit. The dotted lines denote simulations with R = 0 and  $G = \{1, 2, 4\}$ .

Nijjer *et al.* 2018). This rescaling does a reasonable job of collapsing the data for  $G \ll 1$  (figure 6b) and the transition time only weakly depends on G. Note that even when G is small, the initial growth of the spreading length is significantly enhanced when compared with the uniform-density case (figure 6b).

When G is large, the interface slumps on a larger scale which is much faster than the growth of the instability. This form of slumping is analogous to the S-shaped slumping regime in the equal-viscosity case (§ 3.1); however, because the injected fluid is less viscous, the forward-propagating tongue travels faster than the backward-propagating tongue, leading to asymmetric slumping (figure 5b). Based on its similarity with the uniform-viscosity case, we expect that in this regime the spreading length grows like  $h \sim Gt$  and a transition from the diffusive regime at  $t \sim (G^2 P e)^{-1}$ . This rescaling does a reasonable job of collapsing the data for G > O(1) (figure 6c) and the transition time depends only weakly on R.

### 4.2. Intermediate-time gravity-current and fingering regimes

At intermediate times, depending on t and the size of G, there are three possible regimes through which the dynamics evolve. These include a viscous-fingering regime where the interface consists of a set of fine fingers and the spreading length grows linearly in time (figure 4c,i), an injection-driven gravity-current regime where the interface is smooth and the spreading length also grows linearly in time but with a larger prefactor (figure 4d,i), and a density-driven gravity-current regime where the interface is also smooth but the spreading length grows like  $h \sim t^{1/2}$  (not shown). Depending on the magnitude of G only a subset of these regimes are seen. For small G, the interface initially fingers, but over time the fingers coalesce and combine to spread as an injection-driven gravity current. For large G, fingering is suppressed and the fluid spreads as a density-driven gravity current before transitioning to an injection-driven one. In the following subsections we aim to delineate these regimes and develop models for how the transversely averaged concentration evolves in each case.

## 4.2.1. Concentration model

At intermediate times, in all cases, the interface becomes long and thin,  $h \gg 1$  therefore the flow is predominantly horizontal, and longitudinal diffusion is negligible. In this case the flow is in 'vertical flow equilibrium' (Yortsos 1995). Neglecting horizontal gradients in the stream function in (2.21) gives

$$\frac{\mathrm{d}u}{\mathrm{d}y} - R\frac{\partial c}{\partial y}u = R\frac{\partial c}{\partial y} + G\,\mathrm{e}^{Rc}\frac{\partial c}{\partial x}.\tag{4.2}$$

Integrating with respect to y gives

$$u = e^{Rc} \int G \frac{\partial c(x, s, t)}{\partial x} \, \mathrm{d}s + c_1 e^{Rc}, \qquad (4.3)$$

where  $c_1$  is a constant of integration. Integrating again over the transverse direction and imposing no net horizontal flow in the moving frame of reference  $(\int_0^1 u \, dy = 0)$  yields

$$u = \frac{e^{Rc} - \int_0^1 e^{Rc} \, dy}{\int_0^1 e^{Rc} \, dy} + G e^{Rc} \frac{\left(\int_0^1 e^{Rc} \, dy\right) \left(\int \frac{\partial c(x, s, t)}{\partial x} \, ds\right) - \int_0^1 e^{Rc} \left(\int \frac{\partial c(x, s, t)}{\partial x} \, ds\right) dy}{\int_0^1 e^{Rc} \, dy}.$$
 (4.4)

In this limit, there are two main contributions to the horizontal velocity: the first term corresponds to the background pressure gradient due to injection and is driven by the viscosity difference between the two fluids; and the second term corresponds to the horizontal gradient in buoyancy. Multiplying (4.4) by *c* and integrating over the transverse direction gives the advective flux  $\overline{uc'} = \overline{uc}$ . Substituting this flux into (2.17) and neglecting longitudinal diffusion yields a nonlinear advection–diffusion equation for the transversely averaged concentration, as follows:

$$\frac{\partial \bar{c}}{\partial t} + \frac{\partial}{\partial x} \left[ \frac{\int_{0}^{1} c e^{Rc} dy}{\int_{0}^{1} e^{Rc} dy} - \bar{c} + G \frac{\left(\int_{0}^{1} c e^{Rc} dy\right) \left(\int_{0}^{1} c e^{Rc} \left(\int \frac{\partial c(x, s, t)}{\partial x} ds\right) dy\right)}{\int_{0}^{1} e^{Rc} dy} \right]$$
$$- \frac{\partial}{\partial x} \left[ G \frac{\left(\int_{0}^{1} c e^{Rc} dy\right) \left(\int_{0}^{1} e^{Rc} \left(\int \frac{\partial c(x, s, t)}{\partial x} ds\right) dy\right)}{\int_{0}^{1} e^{Rc} dy} \right] = 0.$$
(4.5)

**935** A14-12



Figure 7. Evolution of the transversely averaged concentration (or equivalently the height of the current above the base) found by solving (4.7) for (*a*) small times, (R, G) = (2, 8) and *t* ranging logarithmically from 0.03 to 1 and (*b*) large times, (R, G) = (2, 2) and *t* ranging logarithmically from 1 to 32. The small-time asymptotic limit found by solving (4.8) and large-time asymptotic limit (4.10) are given by dashed black lines.

This equation holds in each of the three identified regimes, since regardless of whether gravity or injection dominates or whether the interface is stable or unstable, the spreading zone always grows longitudinally while remaining transversely finite, leading to a simplification of the velocity (4.4). However, the structure of the concentration field varies and different terms dominate in each regime, resulting in different dynamics.

### 4.2.2. Density-driven and injection-driven gravity-current regimes

When G is sufficiently large, the interface does not finger and little mixing occurs. In this limit we can make the simplifying assumption that the two fluids remain completely segregated; that is, the concentration field is defined by a boundary of height Y(x, t) above the base, such that

$$c = \begin{cases} 1, & 0 \le y \le Y(x, t), \\ 0, & Y(x, t) < y \le 1. \end{cases}$$
(4.6)

Note that by transversely averaging, we find that  $\bar{c} = Y$ . Combining (4.4), (4.5) and (4.6) yields a nonlinear advection–diffusion equation for the evolution of the transversely averaged concentration,

$$\frac{\partial \bar{c}}{\partial t} + \frac{\partial}{\partial x} \left[ \frac{(M-1)\bar{c}(1-\bar{c})}{M\bar{c} + (1-\bar{c})} - \frac{MG\bar{c}(1-\bar{c})}{M\bar{c} + (1-\bar{c})} \frac{\partial \bar{c}}{\partial x} \right] = 0,$$
(4.7)

where  $M = e^{R}$  is the ratio of the ambient to injected viscosity (cf. Pegler *et al.* 2014).

Representative solutions of (4.7) for small and large times are given in figures 7(*a*) and 7(*b*), respectively. In both cases the interface spreads asymmetrically and tends to two different self-similar profiles in time. To understand these two different limits and their transition, we note that the gravitational slumping term, that is, the convective flux that is due to gradients in buoyancy, which is proportional to  $\partial \bar{c}/\partial x$ , is initially very large but decreases over time. This leads to two limiting cases of (4.7). In the small-time limit, or equivalently the large *G* limit, the advective term may be neglected, whereas in the large-time, or equivalently the small *G* limit, the gravitational slumping term to the advective term in (4.7), that is the ratio of the buoyancy contribution to the flux to the viscous contribution, we find that the latter can be neglected when  $-GM(\partial \bar{c}/\partial x)/(M-1) \gg 1$ . For  $M \gg 1$ , this corresponds to  $h \ll G$  and if  $h \sim (Gt)^{1/2}$ , a transition between regimes occurs at  $t \sim O(G)$ .

If  $t \ll O(G)$ , the equations admit a similarity solution of the form  $\bar{c}(\eta_d)$  with similarity variable  $\eta_d = x/(Gt)^{1/2}$  where  $\bar{c}$  satisfies

$$\frac{-\eta_d}{2}\frac{\mathrm{d}\bar{c}}{\mathrm{d}\eta_d} = \frac{\mathrm{d}}{\mathrm{d}\eta_d} \left[ \frac{M\bar{c}(1-\bar{c})}{M\bar{c}+(1-\bar{c})}\frac{\mathrm{d}\bar{c}}{\mathrm{d}\eta_d} \right].$$
(4.8)

The solution of (4.8) with fixed  $\bar{c}$  in the far field is given in figure 7(*a*). The solution has the same similarity variable as that of Huppert & Woods (1995) but a different shape. It is asymmetric and depends on the viscosity ratio of the fluids owing to the fact that the less-viscous injected fluid travels faster than the more viscous ambient fluid. Note that Pegler *et al.* (2014) and Zheng *et al.* (2015) solve the same equation, (4.8), but with different initial and boundary conditions, and find that the spreading length grows like  $h \sim t^{2/3}$ .

If  $t \gg O(G)$ , the buoyancy contribution to the flux can be neglected. In this case the similarity solution, with similarity variable  $\eta_a = x/t$ , satisfies

$$\eta_a \frac{\mathrm{d}\bar{c}}{\mathrm{d}\eta_a} = \frac{\mathrm{d}}{\mathrm{d}\eta_a} \left[ \frac{(M-1)\bar{c}(1-\bar{c})}{M\bar{c}+(1-\bar{c})} \right]. \tag{4.9}$$

Solving, the concentration evolves self-similarly as

$$\bar{c}(x,t) = \begin{cases} 1 & x/t < \frac{1}{M} - 1, \\ \frac{1}{M-1} \left( \sqrt{\frac{M}{x/t+1}} - 1 \right) & \frac{1}{M} - 1 \leqslant x/t \leqslant M - 1, \\ 0 & x/t > M - 1. \end{cases}$$
(4.10)

This is exactly the sharp-interface confined gravity-current solution in Pegler *et al.* (2014) and Zheng *et al.* (2015). This limit is given by the dashed black line in figure 7(*b*). In figure 8(*a*,*b*) we compare the full 2-D numerical simulations with the full solutions of (4.7) as well as the small- and large-time limits of (4.7) for two different values of *G*. When  $t \ll G$  (figure 8*a*, with G = 14 and t = 1), both the sharp-interface model (4.7) and the small-time limit (4.8) give good agreement with the full numerical solutions. When  $t \gg G$  (figure 8*b*, with G = 0.5 and t = 10), the model (4.7) and the large-time limit (4.10) give good agreement with the full solutions in the body of the current; however, the tips tend to propagate slower than predicted. This was also observed in experiments by Pegler *et al.* (2014), who suggested that diffusion was the cause of the slow spreading. In the Appendix (A), we consider the effect of a diffuse region on the propagation of the gravity current. We find that in both the large-time and small-time limits, the diffusive model predicts a much slower tip and gives much better agreement with the 2-D simulations than any of the other models (blue lines in figure 8*a*,*b*).

### 4.2.3. Fingering regime

In § 4.2.2 we considered stable displacements where the two fluids are separated by a nearly sharp interface, which occurred when G was large. However, when G is small the interface can be unstable to viscous fingering. We find that this interface morphology has a pronounced effect on the rate of spreading of the fluids.

In figure 9(*a*), the spreading length for different values of *G* is plotted. When  $G \ge 0.25$  (for the parameters in figure 9), the spreading length grows linearly in time at a rate largely insensitive to *G*. This limit corresponds to the injection-driven gravity-current limit



Figure 8. Plot of the transversely averaged concentration for (*a*) small times, (*R*, *Pe*, *G*, *t*) = (1, 4000, 14, 1) and (*b*) large times (*R*, *Pe*, *G*, *t*) = (2, 4000, 0.5, 10) from the 2-D numerical simulations (black lines). The coloured lines represent the four different model solutions: the full one-dimensional (1-D) sharp-interface model (4.7); the small-time limit of the sharp-interface model (4.8); the large-time limit of the sharp-interface model (4.10); and the diffuse-interface model (A 1) with (4.5). The size of the diffuse region l = 0.03 is chosen to fit the full 2-D numerical simulations.



Figure 9. Plot of h(t) for (R, Pe) = (1, 4000) and different values of G in the intermediate-time regime. (b) Plot of the spreading rate  $\dot{h}$  calculated by least-squares fitting a function of the form  $h = h_0 + \dot{h}t$  to the numerical results for t in the range  $5 \le t \le 10$ , for Pe = 4000 and different G and R. The theoretical predictions for (a) h and (b)  $\dot{h}$  are found from the solution (4.10) with either  $M = e^R$  or  $M = e^R erf(\sqrt{R})/erfi(\sqrt{R})$ , given by dashed and dot–dashed lines, respectively.

discussed in the previous section (dashed black line). When G is small ( $G \leq 0.01$  for the parameters in figure 9a), the interface is unstable to viscous fingering. In this limit, the spreading length also grows linearly in time but at a much slower rate.

To model the evolution of the transversely averaged concentration field in the fingering limit, we start by noting that since G is small, the buoyancy terms in (4.5) can be neglected, resulting in exactly the same model discussed in Nijjer *et al.* (2018) for the uniform-density case, that is

$$\frac{\partial \bar{c}}{\partial t} + \frac{\partial}{\partial x} \left[ \frac{\int_0^1 c\mu(c) \, \mathrm{d}y}{\int_0^1 \mu(c) \, \mathrm{d}y} - \bar{c} \right] = 0.$$
(4.11)

When the flow is unstable, the interface develops a set of fingers which elongate and coarsen. Assuming that the spreading zone is composed of  $n_f(x)$  forward-propagating and  $n_b(x)$  backward-propagating fingers of width  $w_f(x)$  and  $w_b(x)$ , and transverse

viscosity distributions,  $\mu_f = \exp(R(1 - y^2/w_f^2))$  and  $\mu_b = \exp(R(y^2/w_f^2))$ , respectively, (4.11) becomes

$$\frac{\partial \bar{c}}{\partial t} + \frac{\partial}{\partial x} \left( \frac{n_f \int_{-w_f}^{w_f} \exp(R(1 - y^2/w_f^2)) \,\mathrm{d}y}{n_f \int_{-w_f}^{w_f} \exp(R(1 - y^2/w_f^2)) \,\mathrm{d}y + n_b \int_{-w_b}^{w_b} \exp(R(y^2/w_b^2)) \,\mathrm{d}y} - \bar{c} \right) = 0.$$
(4.12)

Applying the areal constraint  $n_f w_f + n_b w_b = 1$  and noting that the total concentration of the forward propagating fingers is  $\bar{c}$ ,  $n_f w_f = \bar{c}$ , the evolution of the transversely averaged concentration is given by

$$\eta_a \frac{\mathrm{d}\bar{c}}{\mathrm{d}\eta_a} = \frac{\mathrm{d}}{\mathrm{d}\eta_a} \left[ \frac{(M_e - 1)\bar{c}(1 - \bar{c})}{M_e \bar{c} + (1 - \bar{c})} \right],\tag{4.13}$$

where  $M_e = e^R \operatorname{erf}(\sqrt{R})/\operatorname{erfi}(\sqrt{R})$ ,  $\operatorname{erf}(x)$  is the error function and  $\operatorname{erfi}(x)$  is the imaginary error function. Note that the average concentration has exactly the same functional form as the injection-driven gravity-current (4.10) but with an effective viscosity contrast  $M_e$ . The starting equation (4.11) is the same in both cases, the only difference between the fingering regime and the injection-driven gravity-current regime is the structure of the concentration field. In the gravity-current limit, the fluids remain mostly segregated, while in the fingering limit, there is significant mixing, resulting in a smaller effective viscosity contrast between the ambient and injected fluids.

Figure 9(*b*) shows the spreading rate, defined as the rate of change of the spreading length,  $\dot{h}$ , as a function of *R* and *G*. For both values of *R*, we find a distinct shift in  $\dot{h}$  from the viscous fingering limit to the gravity current limit. This is in qualitative agreement with the findings of Berg *et al.* (2010), who found an abrupt change in breakthrough times (the time it takes for the injected fluid to transit a fixed length) as *G* was varied. The theoretical predictions for the fingering and injection-driven gravity current regimes are given by dashed and dot-dashed lines, respectively. The fingering limit shows excellent agreement for both values of *R* and the gravity-current limit shows excellent agreement for R = 1, but overestimates the spreading rate for R = 3, which is likely due to mixing at the tip.

In addition to the morphology changing with *G*, it also changes in time. For example, in figure 9(*a*), for G = 0.05, there is an abrupt change in the slope from the fingering limit to the gravity-current limit. This is because as the fingers elongate, they are advected towards the boundaries, coarsening until a single dominant finger remains. This transition occurs at a time  $t \sim O(1/G)$  (the time it takes for the fingers to rise or fall across a length O(1)), and will occur so long as 1/G < O(Pe), that is, this change in morphology occurs before the flow transitions to the late-time regime (§ 4.3).

If G is sufficiently large, such that the fingers coarsen faster than the instability can grow, the instability can be suppressed altogether. Comparing the time scale of the growth of the instability  $(O(1/R^2Pe))$  (Tan & Homsy 1986) with the rise/fall time of the fluid (O(1/G)), suggests that when  $G > O(R^2Pe)$ , the interface will always spread as a gravity current and will not finger. Figure 10 maps the stability boundary in G - R and G - Pephase spaces. We find the transition occurs when  $G \approx 5 \times 10^{-5}R^2Pe$ , in agreement with this simple scaling argument, where the prefactor is determined by fitting the numerical results.



Figure 10. Stable versus unstable displacements for (a) Pe = 4000 and (b) R = 2. Filled circles denote simulations where no fingers were observed during the entire length of the simulations, while unfilled circles denote simulations where fingers were observed for at least some portion of the simulation. Dashed lines show the stability boundary  $G = 5 \times 10^{-5} R^2 Pe$ .



Figure 11. Colourmaps (with overlain contours) of the (a,b) concentration field, (c,d) concentration deviations c' and (e,f) streamwise velocity, u for (a,c,e) (R, Pe, G, t) = (1.5, 1000, 0.1, 1000) (small G) and (b,d,f) (R, Pe, G, t) = (1.5, 100, 10, 1000) (large G). Panels (a,c,e) correspond to flow in the shutdown regime and (b,d,f) correspond to flow in the viscously enhanced Taylor slumping regime. Note that the aspect ratio of the figures is compressed, so variations in the x-direction seem more pronounced than they actually are.

## 4.3. Late-time shutdown and viscously enhanced Taylor slumping regimes

Over long times, diffusion in the transverse direction tends to homogenize the concentration field vertically. The concentration gradient in this case is predominantly in the streamwise direction, as is the fluid flow. In this late-time regime, the interface evolves in two ways. First, in the same manner as the shutdown of the viscous fingering instability, the concentration field is composed of a steady linear background gradient with decaying deviations superimposed. The streamwise velocity closely tracks the deviations and both are horizontally uniform (figure 11a,c,e). Second, the background concentration evolves asymmetrically, with the slope being shallower upstream. The velocity no longer tracks the deviations and neither the velocity nor the concentration is horizontally uniform (figure 11b,d,f).

### 4.3.1. Concentration model

The late-time regime is characterized by a weak background concentration gradient with small deviations superimposed. Assuming the flow is in vertical flow equilibrium, the horizontal velocity is given by (4.4). Decomposing the concentration field into its transverse average and deviations, and making the assumption that  $\partial \bar{c}/\partial x \gg \partial c'/\partial x$ , (4.4) becomes

$$u = \frac{e^{Rc'} - \int_0^1 e^{Rc'} \, dy}{\int_0^1 e^{Rc'} \, dy} + G e^{R\bar{c}} \frac{\partial \bar{c}}{\partial x} \left( \frac{e^{Rc'} \left( \int_0^1 e^{Rc'} \, dy \right) y - e^{Rc'} \int_0^1 e^{Rc'} y \, dy}{\int_0^1 e^{Rc'} \, dy} \right), \quad (4.14)$$

where we have used the fact that  $\int \partial \bar{c} / \partial x \, dy = y \partial \bar{c} / \partial x$ . Next, assuming  $c' \ll 1$  and  $\partial \bar{c} / \partial x \ll 1$  and Taylor expanding, yields

$$u = Rc' + G e^{R\bar{c}} \frac{\partial \bar{c}}{\partial x} \left( y - \frac{1}{2} \right) + O\left( c' \frac{\partial \bar{c}}{\partial x} \right) + O(c'^2).$$
(4.15)

As before, there are two main contributions to the horizontal velocity. The first term, driven by the viscosity difference between the two fluids is, to leading order, proportional to the vertical deviations in the concentration. The second term, driven by the density difference between the two fluids, is only dependent on the longitudinal gradient of the transversely averaged concentration.

By substituting (4.15) into (2.18), and neglecting terms  $O(c^{2})$ , we find that the evolution equation for the deviations is given by

$$\frac{\partial c'}{\partial t} + Rc'\frac{\partial \bar{c}}{\partial x} + Ge^{R\bar{c}} \left(\frac{\partial \bar{c}}{\partial x}\right)^2 \left(y - \frac{1}{2}\right) = \frac{1}{Pe} \frac{\partial^2 c'}{\partial y^2}.$$
(4.16)

Note that at late times, when the spreading length is sufficiently long, concentration is transversely homogenized due to diffusion, faster than it is advected ( $h/u \ll Pe$ ), and leads to a decoupling of (2.18) and (2.17) (cf. Taylor 1953).

Solving (4.16) with no flux boundary conditions in the vertical direction gives

$$c' = \sum_{n \ge 1} \left[ K_n e^{-\gamma_n t} + \frac{2 - 2\cos(\pi n)}{\pi^2 n^2} \frac{G e^{R\bar{c}}}{\gamma_n} \left(\frac{\partial \bar{c}}{\partial x}\right)^2 \right] \cos(\pi n y), \quad (4.17)$$

where  $\gamma_n = R\partial \bar{c}/\partial x + n^2 \pi^2/Pe$ , and  $K_n = 2 \int_0^1 c'(x, y, 0) \cos(\pi ny) dy$  corresponds to the initial conditions at the onset of the regime. When G = 0 this reduces to the late-time, shutdown regime of the miscible viscous-fingering instability (cf. Nijjer *et al.* 2018). When R = 0, this reduces to the Taylor-slumping regime described by Szulczewski & Juanes (2013). In general, when both  $R \neq 0$  and  $G \neq 0$ , the flow transitions from the shutdown regime to a viscously enhanced Taylor-slumping regime.

In the shutdown regime, that is for 'small' t, the exponentially decaying term in (4.17), which is O(1), dominates. The flow is dominated by the slowest decaying mode and c' and u can be approximated as

$$u \approx Rc' \approx RK_1 e^{-\gamma_1 t} \cos\left(\pi y\right). \tag{4.18}$$

The concentration deviations and streamwise velocity are horizontally uniform and decay exponentially. This correspondence between the concentration deviations and the velocity



Figure 12. (a) Evolution of  $\bar{c}$  for (R, Pe, G) = (1.5, 1000, 0.3) and t spaced evenly from 150 to 900. (b) Plot of  $\bar{c}(x)$  at t = 1000 for R ranging from 0.5 to 2.5, G ranging from 0.01 to 0.8 and Pe ranging from 300 to 1000. (c)Plot of Nu(t) for (R, Pe, G) = (1.5, 1000, 0.3). The solid and dashed black lines denote theoretical predictions with K = 0.5 and K = 0.6, respectively.

can be seen in figure 11. Substituting (4.17) into (2.17), we find that the transversely averaged concentration has the steady linear solution

$$\bar{c} = \frac{1}{2} - \alpha x, \tag{4.19}$$

and the fluid steadily fills in the linear profile over time (figure 12*a*). To determine  $\alpha$  and  $\gamma_1$  uniquely we neglect longitudinal diffusion and relate the time-integrated advective flux through the midplane with the net change in concentration of the right-hand half of the domain, namely

$$\int_{0}^{\infty} \int_{0}^{1} uc' \, \mathrm{d}y \, \mathrm{d}t \approx \int_{0}^{\infty} \bar{c}(x) \, \mathrm{d}x = \frac{1}{8\alpha}.$$
(4.20)

Substituting (4.18) into (4.20), and assuming that most of the advective exchange occurs in the shutdown regime,  $\gamma_1 = 2K_1^2 \alpha R$  and so

$$\alpha R = \frac{\pi^2}{(2K_1^2 + 1)Pe}, \quad \gamma_1 = \frac{2K_1^2 \pi^2}{(2K_1^2 + 1)Pe}.$$
(4.21*a*,*b*)

Recall that  $K_1 = 2 \int_0^1 c'(x, y, 0) \cos(\pi y) dy$  is the magnitude of the deviations at the onset of the shutdown regime and is in general unknown. Making the simple assumption that the flow from t = 0 consists of a single sinusoidal mode that spans the width of the channel, such that, the maximum concentration is one and the minimum concentration is 0, then  $c'(x, y, 0) = \cos(\pi y)/2$  and  $K_1 = 1/2$ . However, we find that using this value of  $K_1$  underestimates the length of the spreading zone. By instead fitting  $K_1$  to the numerical simulations, we find  $K_1 \simeq 0.6$ . This slightly larger value of  $K_1$  accounts for nonlinear effects as well as the fact that the deviations are not sinusoidal and consist of many modes from t = 0. With this value for  $K_1$  we find much better agreement with the numerical simulations for a wide range of simulations (figure 12b,c).

In the viscously enhanced Taylor-slumping regime, i.e. for large t, the exponentially decaying terms in (4.17) and (4.15) become negligible and the gravitational



Figure 13. (a) The similarity solution of (4.25) for R = 0, 0.5, 1, 1.5, 2, 2.5, 3. The analytical solution for R = 0 is given by the black line (Szulczewski & Juanes 2013). (b) The evolution of  $\bar{c}$  for (R, Pe, G) = (3, 10, 10) at  $t = \{1000, 2000, 4000\}$ . The theoretical predictions, found by solving (4.24), are denoted by dotted lines.

terms dominate. Expanding c' and u in powers of  $\partial \bar{c} / \partial x$  we find that

$$c' = G e^{R\bar{c}} Pe\left(\frac{\partial \bar{c}}{\partial x}\right)^2 \left(\frac{-1}{24} + \frac{y^2}{4} - \frac{y^3}{6}\right) + O\left(\left(\frac{\partial \bar{c}}{\partial x}\right)^3\right),$$
  
$$u = G e^{R\bar{c}} \left(\frac{\partial \bar{c}}{\partial x}\right) \left(y - \frac{1}{2}\right) + O\left(\left(\frac{\partial \bar{c}}{\partial x}\right)^2\right).$$
(4.22)

Combining these expressions, the advective flux is

$$\int_{0}^{1} uc' \, \mathrm{d}y = \frac{PeG^2 \mathrm{e}^{2R\bar{c}}}{120} \left(\frac{\partial\bar{c}}{\partial x}\right)^3 + O\left(\left(\frac{\partial\bar{c}}{\partial x}\right)^5\right). \tag{4.23}$$

Substituting this flux into (2.17) yields a nonlinear diffusion equation for the evolution of  $\bar{c}$ ,

$$\frac{\partial \bar{c}}{\partial t} = \frac{\partial}{\partial x} \left[ \left( \frac{1}{Pe} + \frac{Pe}{120} \left( G e^{R\bar{c}} \frac{\partial \bar{c}}{\partial x} \right)^2 \right) \frac{\partial \bar{c}}{\partial x} \right].$$
(4.24)

By neglecting the effects of molecular diffusion, (4.24) admits a similarity solution of the form  $\bar{c}(\eta)$  with similarity variable  $\eta = x/(G^2 Pe t/120)^{1/4}$ , where  $\bar{c}$  satisfies the nonlinear differential equation

$$\frac{\mathrm{d}}{\mathrm{d}\eta} \left[ \mathrm{e}^{2R\bar{c}} \left( \frac{\mathrm{d}\bar{c}}{\mathrm{d}\eta} \right)^3 \right] = -\frac{\eta}{4} \frac{\mathrm{d}\bar{c}}{\mathrm{d}\eta}.$$
(4.25)

The solution of (4.25) for different values of *R* is given in figure 13(*a*). When R = 0, the interface spreads symmetrically, whereas when R > 0, the interface slumps preferentially upstream.

The solution to the full diffusion equation (4.24) along with the full numerical simulations are plotted in figure 13(b). We find very good agreement between the numerical simulations and the reduced model. Note that for the parameters in figure 13(b), the similarity solution is not able to completely reproduce the full numerical simulations because the effects of molecular diffusion are not negligible. For sufficiently large *Pe*, molecular diffusion can be neglected; however, these sets of parameters are currently beyond the scope of our numerical simulations.



Figure 14. Representative plots of the scaling exponent of the spreading length,  $\delta$ , found by locally fitting a power law of the form  $h = At^{\delta}$ , for R = 1.5 and (a) Pe = 100, (b) Pe = 1000. The regime boundaries (black lines) divide the (I) early-time diffusive, (II) large-G slumping, (III) small-G slumping and viscous fingering, (IV) density-driven gravity current, (V) injection-driven gravity current, (VI) shutdown, (VII) central-finger and boundary-finger and (VIII) viscously enhanced Taylor-slumping regimes. Over long times, the interface evolves through longitudinal diffusion (regime IX not shown).

The transition between the shutdown regime and the viscously enhanced Taylorslumping regime occurs when the two limiting solutions for  $\bar{c}$ , (4.19) and the solution of (4.25), overlap, that is at  $t \sim R^4 P e^3/G^2$ . In fact, if G is sufficiently large such that the gravitational term in (4.17) is O(1), the shutdown regime may be skipped entirely, that is, when  $G \sim O(R^2 P e)$ . Note that this is the same scaling as the transition from stable to unstable displacements described in § 4.2.3, but with a larger prefactor.

Finally, we end this discussion of the late-time dynamics by noting that over very long times, the viscously enhanced Taylor-slumping term in (4.24) becomes negligible compared with the molecular diffusion term, occurring at a time  $t \sim O(G^2 P e^3)$ , and the interface evolves through longitudinal dispersion again.

#### 5. Discussion and conclusions

## 5.1. Summary

In this work, the range of dynamics that are possible when a more viscous fluid is displaced by a less-viscous fluid of a different density during horizontal miscible displacements, with gravity acting perpendicular to the prevailing flow, are examined. Figure 14 delineates the different possible regimes in G - t phase space for two different values of *Pe*. In each case the instantaneous scaling exponent of the spreading length  $\delta$ , where  $h = At^{\delta}$ , is plotted for representative simulations. In general, nine different regimes are possible. The limit  $G \gg 1$  (and  $R \sim O(1)$ ) corresponds to the limit where buoyancy dominates over viscous effects and the limit  $G \ll 1$  (and  $R \sim O(1)$ ) corresponds to the limit where buoyancy is negligible in comparison with viscous effects.

At very early times (regime I), the interface is very sharp and diffusion across it dominates leading to  $t^{1/2}$  growth of the interface. Once advection outpaces diffusion the interface begins to slump (regimes II,III; § 4.1). The dynamics are either dominated by buoyancy differences leading to a transition at  $t \sim O(1/G^2Pe)$ , or by viscosity differences leading to a transition at  $t \sim O(1/R^2Pe)$ . In either case, the interface is sharp and vertical flow is important. Over time, the interface becomes long and thin and vertical flow becomes unimportant. If fingers coarsen due to gravity faster than they can initially grow

 $(G > O(R^2 Pe))$  the interface is stabilized and does not finger. In this case, spreading is initially dominated by gravitational slumping where the rate of spreading diminishes with the buoyancy gradient leading to sublinear growth of the spreading zone  $h \sim t^{1/2}$ (regime IV; § 4.2.1). After  $t \sim O(G)$  the spreading rate becomes dominated by viscous contributions arising from the background pressure gradient (regime V; 4.2.1) leading to linear growth of the spreading zone. If, however,  $G < O(R^2 Pe)$ , the interface can finger, which on average leads to linear growth of the spreading length in time (regime III; 4.2.1). Note that, because the dynamics are chaotic, the scaling exponent varies over time (note the speckle in figure 14), but, on average, the interface spreads linearly. If G < O(Pe), the fingers coalesce until a single dominant finger is left propagating in the centre of the domain with counter-propagating fingers along the top and bottom boundaries. As the single-finger exchange-flow decays, a pair of wider counter-propagating fingers propagating along the boundaries manifest themselves (figure 3e, f), which also propagate and slow down leaving a well-mixed interior (regime VII;  $\S$  3.2). If G > O(Pe), either the interface is stable or it fingers and the fingers coarsen along the boundaries. Eventually, after  $t \sim O(Pe)$ , diffusion homogenizes the concentration field transversely and the shutdown regime is reached (regime VI; §4.3). In this regime, the fluid flow slows over time as the spreading rate, proportional to the transverse variations in viscosity, diffuse away leaving a linear background concentration gradient that is filled in exponentially (i.e.  $\delta \to 0$ ). After  $t \sim O(R^4 P e^3/G^2)$ , the density difference between the two fluids becomes important again and the interface evolves through viscously enhanced Taylor-slumping (regime VIII;  $\S4.3$ ). Over very long times, as the interface becomes more diffuse, Taylor-slumping becomes negligible and the interface evolves through longitudinal diffusion again with the same solution as in regime I (regime IX; not shown).

## 5.2. Implications for geological carbon sequestration

To highlight the different regimes through which the flow can evolve, given physically motivated parameters, we consider the sequestration of CO<sub>2</sub> at Sleipner (Bickle et al. 2007; Boait et al. 2012), In Salah (Vasco et al. 2010) and Salt Creek (Bickle et al. 2017). For illustrative purposes, we make a series of simplifying assumptions, namely that the porous medium is 2-D and homogeneous even though the injection sites have complex geometries and are heterogeneous, and that the two fluids are fully miscible, even though the injected and ambient fluids are only partially miscible and the miscibility varies across the three different scenarios. We take the properties of the fluids to be: viscosity of the injected CO<sub>2</sub>,  $\mu_1 = 6 \times 10^{-5}$  Pa s; viscosity of the ambient fluid,  $\mu_2 =$  $7 \times 10^{-4}$  Pa s; density of the injected CO<sub>2</sub>,  $\rho_1 = 7 \times 10^2$  kg m<sup>-3</sup>; density of the ambient fluid (brine),  $\rho_2 = 1 \times 10^3$  kg m<sup>-3</sup> (Huppert & Neufeld 2014); diffusion coefficient  $D = 2 \times 10^{-9} \text{ m}^2 \text{ s}^{-1}$  (Cadogan, Maitland & Trusler 2014). We make the simplifying assumption that the diffusivity is a constant and equal to the molecular diffusivity of carbon dioxide in brine. This is reasonable as long as the pore-scale Péclet number  $Pe_p$ , defined in § 2, remains small. Simple scaling estimates of  $Pe_p$  indicate that this is the case at In Salah, although Sleipner and Salt Creek may be closer to the limit of validity of this assumption. Furthermore, although these properties will vary significantly between these three case studies given the different depths, temperatures and ambient fluid compositions (at Salt Creek and In Salah, CO<sub>2</sub> is injected into depleted oil fields, whereas at Sleipner CO<sub>2</sub> is injected into a saline aquifer), for simplicity we will assume they are the same in all three cases. We assume that the main differences between the three injection scenarios are the injection rates, permeabilities and thicknesses of the formations.

	Sleipner	In Salah	Salt Creek
$t_{dim} = \frac{a^2\phi}{Q}$	$3.5 \times 10^6 \text{ s}$	$6  imes 10^6  ext{ s}$	$7  imes 10^3  ext{ s}$
$G = \frac{\Delta \rho g k a}{Q \mu_2}$	5	0.08	0.02
$Pe = \frac{Q}{D}$	$2 \times 10^4$	$5 \times 10^3$	$1.5  imes 10^4$
$R = \log\left(\frac{\mu_2}{\mu_1}\right)$	2.5	2.5	2.5

Table 1. Characteristic advective time scale  $t_{dim}$  and dimensionless variables G, R, Pe for the three carbon dioxide sequestration case studies.

The Utsira formation at Sleipner is 200 m thick and formed of nine distinct layers and so we take our representative length scale to be thickness of one layer or a = 20 m. The formation is relatively homogeneous and has a permeability of  $k = 2.5 \times 10^{-12}$  m<sup>2</sup> and porosity  $\phi = 0.35$ . Carbon dioxide is injected at a rate of approximately  $Q = 4 \times 10^{-5}$  m<sup>2</sup> s<sup>-1</sup>. The Krechba formation at In Salah has a similar characteristic length scale a = 20 m, however, it is less porous,  $\phi = 0.15$ , less permeable,  $k = 1 \times 10^{-14}$  m<sup>2</sup>, and injection is slower,  $Q = 1 \times 10^{-5}$  m<sup>2</sup> s<sup>-1</sup>. The Frontier formation at Salt Creek is 20 m thick and highly heterogeneous, and fluid flow is believed to be dominated by a few layers that are 1 m thick with  $\phi = 0.2$  and  $k = 1.5 \times 10^{-13}$  m<sup>2</sup>. We estimate the injection velocity by dividing the distance between the injection and production wells by the breakthrough time of the bulk of the CO<sub>2</sub>,  $Q = 3 \times 10^{-5}$  m<sup>2</sup> s<sup>-1</sup>.

The relevant non-dimensional parameters for these three case studies are summarized in table 1. Comparing the three scenarios, we find gravity to be relatively important at Sleipner but unimportant at In Salah or Salt Creek. Comparing G with the critical  $G_{crit}$ for fingering, we expect that Sleipner is mostly stable to fingering ( $G_{crit} \approx 6$ ), whereas at In Salah ( $G_{crit} \approx 1$ ) and Salt Creek ( $G_{crit} \approx 5$ ) we expect that the interface is initially dominated by fingering. We summarize the dimensional time scales for the different regimes in each of the three cases in figure 15. Under our assumptions, we find that the dynamics at Sleipner are currently dominated by the large-time gravity current limit, the dynamics at In Salah, by the end of injections, were dominated by viscous fingering and the large-time gravity current limits, while at Salt Creek, at the time of breakthrough, we expect that the dynamics were dominated by the shutdown regime. Although our model accounts for density and viscosity differences, a number of simplifying assumptions were made that may not hold in these settings. For instance, the ambient and injected fluids were assumed to be fully miscible, even though in reality they are only partially miscible, likely resulting in an overestimation of dissolution rates. The diffusivity was also assumed to be constant and equal to the molecular diffusivity of carbon dioxide in brine, even though mechanical dispersion may not be negligible in these three scenarios, and the porous media were assumed to be homogeneous, even though geological formations tend to be heterogeneous on a range of scales from the pore scale to the reservoir scale. These latter two assumptions likely result in an underestimation of dispersion. While some of these factors have been studied in isolation (Golding et al. 2011; Nicolaides et al. 2015; Fu, Cueto-Felgueroso & Juanes 2018; Nijjer et al. 2019), how these different factors combine to effect spreading in porous media remains to be understood.



Figure 15. Evolution of the displacement front in the three case studies. The black dots denote the time since injection at Sleipner, the total injection time at In Salah, and time until breakthrough at Salt Creek.

Finally, one key assumption made in this work is that the flow within the porous media may be treated as 2-D. In reality, in the case studies considered, and in porous media in general, geometries are complex and three-dimensional, for which the exact details of each regime, and the transitions between them, need to be determined. Nonetheless, we expect that many of our analytical results can readily extended to three dimensions. For instance, the intermediate-time gravity-current, late-time shutdown and late-time viscously enhanced Taylor slumping regimes can be extended to three dimensions by considering a finite third dimension or axisymmetry, i.e. flow from a localized source. While the microscopic dynamics of the unstable fingering process may differ in three dimensions from the 2-D case considered here, the overall spreading behaviour has generally been found to be insensitive of the dimensionality of the nonlinear fingering process (e.g. Zimmerman & Homsy 1992; Tchelepi & Orr 1994).

In summary, we have investigated the combined effects of viscous fingering and gravitational segregation in miscible displacements through porous media. In doing so, we have identified the different possible regimes through which the flow evolves and demarcated the boundaries between these different regimes. In addition to being relevant to the understanding of geological storage of carbon dioxide, these results are applicable in a wide range of contexts including in enhanced oil recovery (Lake 1989), mantle dynamics (Schoonman, White & Pritchard 2017), contaminant transport (Abriola 1987), food processing (Hill 1952) and chromatography (Catchpoole *et al.* 2006).

**Funding.** J.S.N. was supported by the Bill and Melinda Gates Foundation [OPP1144] and J.A.N. acknowledges support from the GeoCquest consortium, a BHP-supported collaborative project between Cambridge, Stanford and Melbourne Universities.

Declaration of interests. The authors report no conflict of interest.

#### Author ORCIDs.

- Japinder S. Nijjer https://orcid.org/0000-0002-0972-8095;
- Duncan R. Hewitt https://orcid.org/0000-0001-6190-5514;
- Jerome A. Neufeld https://orcid.org/0000-0002-3284-5169.

## Appendix A. The effect of mixing on current propagation

In § 4.2.1 we assumed a sharp interface separated the two fluids. Here we consider the effect of a diffuse interface on its evolution. In a given vertical slice, for a given transversely averaged concentration  $\bar{c}$ , the relative flow of the two fluids is maximized when there is a binary distribution of concentration, and therefore mixing in general tends to slow their relative velocities.

To identify the effect of diffusion and mixing on the shape and evolution of the transversely averaged concentration field, we make the ansatz that

$$c = \frac{1}{2} + \frac{1}{2} \operatorname{erf}\left(\frac{\bar{c}(x,t) - y}{l}\right),$$
 (A1)

instead of (4.6), where l is the width of the diffuse region, which we expect depends on the parameters in the problem. Substituting (A1) into (4.5), results in a nonlinear advection–diffusion equation with one additional parameter l. With this model we now include the effects of a diffuse boundary layer on the propagation of the gravity current; however, we ignore any spatial variations in the thickness of the boundary layer and any time dependence. We make this simplification assuming that the steepening of the concentration gradient due to stretching at the interface balances diffusion leading to a boundary layer that only varies slowly (see, for e.g. de Anna *et al.* 2014).

#### REFERENCES

- ABDUL HAMID, S.A. & MUGGERIDGE, A.H. 2020 Fingering regimes in unstable miscible displacements. *Phys. Fluids* **32** (1), 016601.
- ABRIOLA, L.M. 1987 Modeling contaminant transport in the subsurface: an interdisciplinary challenge. *Rev. Geophys.* **25** (2), 125–134.
- ADAMS, J.C. 1993 MUDPACK-2: Multigrid software for approximating elliptic partial-differential equations on uniform grids with any resolution. *Appl. Math. Comput.* 53, 235–249.
- DE ANNA, P., DENTZ, M., TARTAKOVSKY, A. & LE BORGNE, T. 2014 The filamentary structure of mixing fronts and its control on reaction kinetics in porous media flows. *Geophys. Res. Lett.* 41, 4586–4593.
- BERG, S., OEDAI, S., LANDMAN, A.J., BRUSSEE, N., BOELE, M., VALDEZ, R. & VAN GELDER, K. 2010 Miscible displacement of oils by carbon disulfide in porous media: experiments and analysis. *Phys. Fluids* 22, 113102.
- BICKLE, M.J., CHADWICK, R.A., HUPPERT, H.E., HALLWORTH, M.A. & LYLE, S. 2007 Modelling carbon dioxide accumulation at Sleipner: implications for underground carbon storage. *Earth Planet. Sci. Lett.* 255, 164–176.
- BICKLE, M., KAMPMAN, N., CHAPMAN, H., BALLENTINE, C., DUBACQ, B., GALY, A., SIRIKITPUTTISAK, T., WARR, O., WIGLEY, M. & ZHOU, Z. 2017 Rapid reactions between CO<sub>2</sub>, brine and silicate minerals during geological carbon storage: modelling based on a field CO<sub>2</sub> injection experiment. *Chem. Geo.* 468, 17–31.
- BOAIT, F.C., WHITE, N.J., BICKLE, M.J., CHADWICK, R.A., NEUFELD, J.A. & HUPPERT, H.E. 2012 Spatial and temporal evolution of injected CO<sub>2</sub> at the Sleipner Field, North Sea. J. Geophys. Res. 117, B03309.
- CADOGAN, S.P., MAITLAND, G.C. & TRUSLER, J.P.M. 2014 Diffusion coefficients of CO<sub>2</sub> and N<sub>2</sub> in water at temperatures between 298.15 K and 423.15 K at pressures up to 45 MPa. *J. Chem. Engng Data* **59**, 519–525.
- CAMHI, E., MEIBURG, E. & RUITH, M. 2000 Miscible rectilinear displacements with gravity override. Part 2. Heterogeneous porous media. J. Fluid Mech. 420, 259–276.
- CATCHPOOLE, H.J., SHALLIKER, R.A., DENNIS, G.R. & GUIOCHON, G. 2006 Visualising the onset of viscous fingering in chromatography columns. J. Chromatogr. 1117, 137–145.
- FLEURY, P., BAKALOWICZ, M. & DE MARSILY, G. 2007 Submarine springs and coastal karst aquifers: a review. J. Hydrol. 339, 79–92.
- FU, X., CUETO-FELGUEROSO, L. & JUANES, R. 2018 Nonequilibrium thermodynamics of hydrate growth on a gas-liquid interface. *Phys. Rev. Lett.* **120**, 144501.

- GOLDING, M.J., NEUFELD, J.A., HESSE, M.A. & HUPPERT, H.E. 2011 Two-phase gravity currents in porous media. J. Fluid Mech. 678, 248–270.
- HESSE, M.A., TCHELEPI, H.A., CANTWELL, B.J. & ORR, F.M. JR. 2007 Gravity currents in horizontal porous layers: transition from early to late self-similarity. *J. Fluid Mech.* **577**, 363–383.
- HILL, S. 1952 Channelling in packed columns. Chem. Engng Sci. 1, 247–253.
- HOMSY, G.M. 1987 Viscous fingering in porous media. Annu. Rev. Fluid Mech. 19, 271-311.
- HUPPERT, H.E. & NEUFELD, J.A. 2014 The fluid mechanics of carbon dioxide sequestration. Annu. Rev. Fluid Mech. 46, 255–272.
- HUPPERT, H.E. & WOODS, A.W. 1995 Gravity-driven flows in porous layers. J. Fluid Mech. 292, 55-69.
- LAKE, L.W. 1989 Enhanced Oil Recovery. Prentice Hall.
- NICOLAIDES, C., JHA, B., CUETO-FELGUEROSO, L. & JUANES, R. 2015 Impact of viscous fingering and permeability heterogeneity on fluid mixing in porous media. *Water Resour. Res.* **51**, 2634–2647.
- NIJJER, J.S., HEWITT, D.R. & NEUFELD, J.A. 2018 The dynamics of miscible viscous fingering from onset to shutdown. J. Fluid Mech. 837, 520–545.
- NIJJER, J.S., HEWITT, D.R. & NEUFELD, J.A. 2019 Stable and unstable miscible displacements in layered porous media. J. Fluid Mech. 869, 468–499.
- PEGLER, S.S., HUPPERT, H.E. & NEUFELD, J.A. 2014 Fluid injection into a confined porous layer. J. Fluid Mech. 745, 592–620.
- RIAZ, A. & MEIBURG, E. 2003 Three-dimensional miscible displacement simulations in homogeneous porous media with gravity override. J. Fluid Mech. 494, 95–117.
- ROGERSON, A. & MEIBURG, E. 1993a Numerical simulation of miscible displacement processes in porous media flows under gravity. *Phys. Fluids* 5, 2644–2660.
- ROGERSON, A. & MEIBURG, E. 1993b Shear stabilization of miscible displacement processes in porous media. *Phys. Fluids* 5, 1344–1355.
- RUITH, M. & MEIBURG, E. 2000 Miscible rectilinear displacements with gravity override. Part 1. Homogeneous porous medium. J. Fluid Mech. 420, 225–257.
- SCHOONMAN, C.M., WHITE, N.J. & PRITCHARD, D. 2017 Radial viscous fingering of hot asthenosphere within the Icelandic plume beneath the North Atlantic Ocean. *Earth Planet. Sci. Lett.* 468, 51–61.
- SIMMONS, C.T., FENSTEMAKER, T.R. & SHARP, J.M. JR. 2001 Variable-density groundwater flow and solute transport in heterogeneous porous media: approaches, resolutions and future challenges. J. Contam. Hydrol. 52, 245–275.
- SUEKANE, T., KOE, T. & BARBANCHO, P.M. 2019 Three-dimensional interaction of viscous fingering and gravitational segregation in porous media. *Fluids* 4, 130.
- SZULCZEWSKI, M.L. & JUANES, R. 2013 The evolution of miscible gravity currents in horizontal porous layers. J. Fluid Mech. 719, 82–96.
- TAN, C.T. & HOMSY, G.M. 1986 Stability of miscible displacements in porous media: rectilinear flow. *Phys. Fluids* 29, 3549–3556.
- TAN, C.T. & HOMSY, G.M. 1988 Simulation of nonlinear viscous fingering in miscible displacement. *Phys. Fluids* 31, 1330–1338.
- TAYLOR, G.I. 1953 Dispersion of soluble matter in solvent flowing slowly through a tube. *Proc. R. Soc. Lond.* A **219** (1137), 186–203.
- TCHELEPI, H.A., ORR, F.M., RAKOTOMALALA, N., SALIN, D. & WOUMENI, R. 2004 Stabilizing viscosity contrast effect on miscible displacement in heterogeneous porous media, using lattice Bhatnagar–Gross–Krook simulations. *Phy. Fluids* 16 (12), 4408–4411.
- TCHELEPI, H.A. & ORR, F.M. JR. 1994 Interaction of viscous fingering, permeability heterogeneity, and gravity segregation in three dimensions. *SPE Res. Eng.* 9, 266–271.
- VASCO, D.W., RUCCI, A., FERRETTI, A., NOVALI, F., BISSELL, R.C., RINGROSE, P.S., MATHIESON, A.S. & WRIGHT, I.W. 2010 Satellite-based measurements of surface deformation reveal fluid flow associated with the geological storage of carbon dioxide. *Geophys. Res. Lett.* 37, L03303.
- WOODS, A.W. 1999 Liquid and vapor flow in superheated rock. Annu. Rev. Fluid Mech. 31, 171-199.
- YORTSOS, Y.C. 1995 A theoretical analysis of vertical flow equilibrium. Trans. Porous Med. 18, 107-129.
- ZHENG, Z., GUO, B., CHRISTOV, I.C., CELIA, M.A. & STONE, H.A. 2015 Flow regimes for fluid injection into a confined porous medium. J. Fluid Mech. 767, 881–909.
- ZIMMERMAN, W.B. & HOMSY, G.M. 1992 Three dimensional viscous fingering: a numerical study. *Phys. Fluids* 4, 1901–1914.