

LUGOVSKOE, WESTERN SIBERIA: A POSSIBLE EXTRA-ARCTIC MAMMOTH REFUGIUM AT THE END OF THE LATE GLACIAL

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ABSTRACT. Eleven woolly mammoth bone samples from Lugovskoe (central West Siberian Plain, Russia) were radiocarbon dated in 3 laboratories: Institute of Geology, Novosibirsk; Oxford University, Oxford; and Christian Albrechts University, Kiel. Each laboratory used its own protocol for collagen extraction. Parallel dating was carried out on 3 samples in Novosibirsk and Oxford. Two results are in good agreement. However, there is a major discrepancy between 2 dates obtained for the third sample. The dates obtained so far on the Lugovskoe mammoths range from about 18,250 BP to about 10,210 BP. The Lugovskoe results thus far confirm the possibility of woolly mammoth survival south of Arctic Siberia in the Late Glacial after about 12,000 BP, which has important implications for interpreting the process of mammoth extinction. The site has also produced the first reliable traces of human occupation from central Western Siberia at the Late Glacial, including unique direct evidence of mammoth hunting.

INTRODUCTION

Establishing a chronology of Pleistocene megafaunal extinctions is very important for understanding their cause(s), and placing the events in the wider context of Quaternary paleoenvironment history (Martin and Klein 1984; Stuart 1991; MacPhee 1999). A large radiocarbon data set is available for the woolly mammoth (*Mammuthus primigenius* Blum; hereafter, “mammoth”) (Stuart 1991; Sulerzhitsky 1997; Vasil’chuk et al. 1997; Kuzmin et al. 2003). Previous research suggests that mammoths retreated at about 12,000 BP from most of its former extensive distribution in Northern Eurasia, but survived in the far north of mainland Siberia for another 2 millennia (cf. Sulerzhitsky 1997; Sher 1997). This model of spatial-temporal pattern of mammoth dynamics at the end of the Pleistocene can be called the “retreat to the north after ca. 12,500–12,000 BP” (cf. Sher 1997).

Recent research, however, indicates that mammoths were present until about 10,000 BP in the Baltic States and North Russian Plain, outside of the Arctic regions (Stuart et al. 2002: 1567; Lõugas et al. 2002). These European occurrences probably represent recolonization from the refugium in Northern Siberia during the renewed cold of the Younger Dryas phase (Stuart, forthcoming). The possibility that mammoths also survived in Siberia outside the Arctic region after 12,000 BP is very important in interpreting the process of mammoth extinction. Lugovskoe, in the central West Siberian Plain, is a key locality for the study of mammoth’s late survival because it was dated to the critical period of about 12,000–10,300 BP. ¹⁴C dates on woolly mammoth remains were carried out in 2000–2002. Here, we present the ¹⁴C results obtained so far and their preliminary interpretation.

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MATERIAL AND METHODS

The Pleistocene fossil mammal locality of Lugovskoe is in the modern forest zone of the central Western Siberian Plain (60°57'N; 68°32'E) about 25 km west of Khanty-Mansiysk City and the confluence of the Ob and Irtysh Rivers (Figure 1). Paleontological research has been conducted here since 1994 (Mashchenko et al. 2003). The bone bed occurs at a depth of 0.5–1.5 m below the modern surface within the muddy clay deposits of a small stream that flows into the Ob River. Although it is impossible to conduct proper excavations because the deposits are saturated with stream water, about 4500 fossil mammal bones were recovered in 1998–2002, of which 98% belong to mammoth. Remains of at least 27 individual mammoths were collected from the muddy clay, and several skeletal elements were found in anatomical association (Pavlov et al. 2002). Lugovskoe represents one of the richest so-called “mammoth cemeteries” in Siberia.

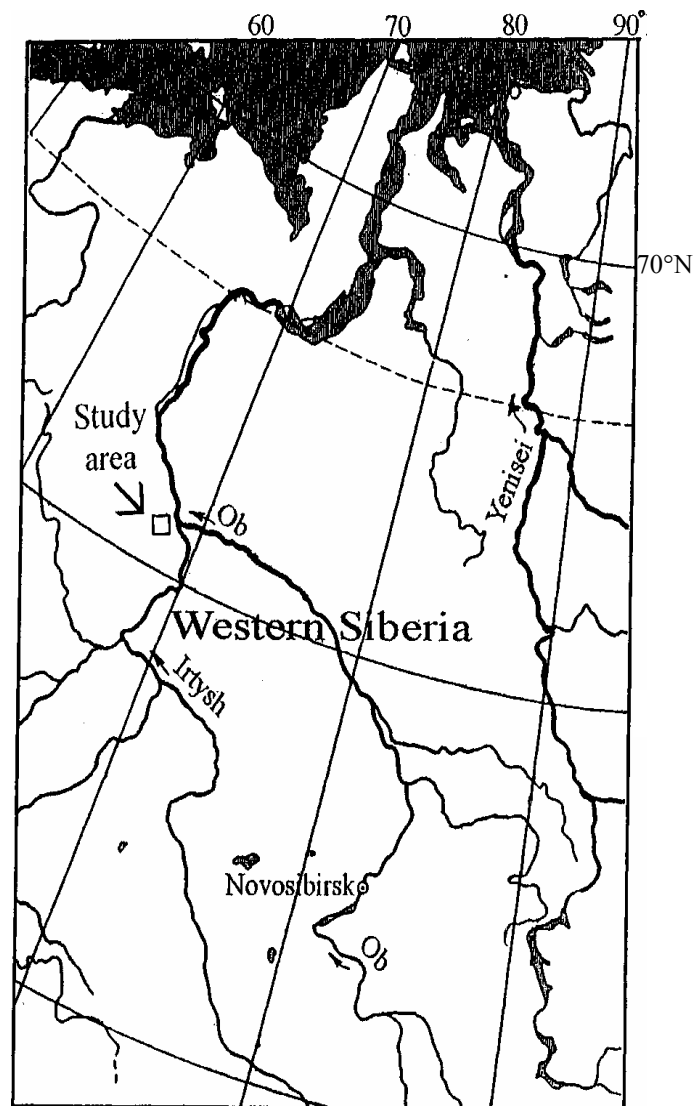


Figure 1 The position of the Lugovskoe locality in Western Siberia

^{14}C dating of the mammoth bones was undertaken in 3 laboratories: the Institute of Geology, Novosibirsk (Russia); Oxford University, Oxford (UK); and Christian Albrechts University, Kiel (Germany). Collagen extracted from mammoth bones was dated by the liquid scintillation counting (LSC) (Novosibirsk) and accelerator mass spectrometry (AMS) (Oxford and Kiel) techniques. In each laboratory, the collagen extraction protocol is different.

In Novosibirsk, pieces of bone 10–20 cm long were mechanically cleaned of any surface contamination and were demineralized with a 5% HCl solution (at 7–8 L of solution for 1 kg of bone) at 2–3 °C. Since the process may take 1–2 weeks, the sample is refrigerated. The extracted gelatin-like collagen was thoroughly washed with distilled water. The collagen was then treated with an 0.1-N solution of sodium base (NaOH) for several hours in order to remove the humic acids. The remaining collagen was again washed with distilled water, dried, and carbonized by heating in an 800 °C oxygen-free environment. To remove the phosphorous compounds, the carbonized collagen was treated with a mixture of nitric acid (HNO_3) and HCl (“aqua regia”). Finally, the purified collagen was washed with distilled water, dried, and used for benzene preparation and LSC measurements.

At the Oxford Radiocarbon Accelerator Unit (ORAU), bones were cleaned by scraping with a scalpel to remove soil coloration, sediment detritus, and weathered surfaces. Bone was then chemically pretreated using the ORAU continuous-flow semi-automated method (Bronk Ramsey et al., these proceedings), with an acid-base-acid sequence to isolate crude collagen, then gelatinization with weakly acidic water (pH=3) for 20 hr in an incubator at 75 °C. Ultrafilters were used to isolate the >30-kD fraction of the bone hydrolyzate. This fraction was then lyophilized. Samples were combusted using a Europa Scientific ANCA-MS system consisting of a 20-20 IR mass spectrometer interfaced to a Roboprep CHN sample converter unit operating in continuous flow mode. Graphite was prepared by reduction of CO_2 over an iron catalyst in an excess H_2 atmosphere at 560 °C prior to AMS ^{14}C measurement (Bronk Ramsey and Hedges 1999; Bronk Ramsey et al. 2000).

At the Leibnitz Laboratory for Radiometric Dating and Isotope Research (Kiel), bone sample was checked and mechanically cleaned under the microscope. The collagen content was calculated from the concentration of nitrate, determined by colorimetry as nitrate in about 2 mg of bone material. The sample was first treated with acetone, rinsed with demineralized water, and subsequently demineralized in 1% HCl. This was followed by 1% NaOH at room temperature for 1 hr to remove humic acids, then washed and subsequently acidified with 1% HCl. After filtering, the purified collagen was gelatinized overnight in 1.6 mL of H_2O at 90 °C and pH=3. The gelatin solution was dried in quartz tubes, to which CuO and silver wool was added and combusted to CO_2 at 900 °C. The CO_2 gas was reduced with H_2 over about 2 mg of Fe powder as catalyst, and the resulting carbon/iron mixture was pressed into a pellet for AMS analysis.

RESULTS AND DISCUSSION

Measurements of stable carbon and nitrogen isotopes ($\delta^{13}\text{C}$ and $\delta^{15}\text{N}$) were made at the Oxford Laboratory for 3 samples, OxA-12029, 12030, and 12031. $\delta^{13}\text{C}$ values are within –20.7 to –21.3‰ (Table 1), and $\delta^{15}\text{N}$ values are within +4.7 to +5.9‰. The collagen yield from 3 samples dated at Oxford is 1.9–4.8%, which indicates poor preservation and low survival of collagen >30 kD MW. The collagen yield for 1 sample dated at Kiel (KIA-19643) was 23%, about the content of a fresh bone, indicating excellent preservation. The minimum threshold for yield acceptance at Oxford is 10 mg/g^{-1} (1% weight collagen). Differences between the collagen yields by laboratory are liable to be a function of pretreatment method. The Oxford ultrafiltration method often produces lower yields

in comparison with simpler bone hydrolyzate filtration methods, a fact we attribute to the removal of degraded collagen, salts, and low molecular weight particulates. The C/N ratios for the Oxford samples ranged between 3.3–3.4, and values between 2.9–3.5 are considered acceptable (DeNiro 1985).

Table 1 ^{14}C dates for woolly mammoth from Lugovskoe^a.

Lab code	Sample	^{14}C age (BP)	$\delta^{13}\text{C}$ (‰)	$\delta^{15}\text{N}$ (‰)	C/N ratio
SOAN-3838	Tibia	18,250 ± 1100	—	—	—
SOAN-5065	Scapula fragment	15,420 ± 215	—	—	—
SOAN-4940	Thoracic vertebra ^b	13,720 ± 160	—	—	—
OxA-12030	Thoracic vertebra ^b	13,455 ± 60	-21.1	+4.7	3.4
SOAN-4942	Pelvic bone ^c	13,490 ± 155	—	—	—
OxA-12029	Pelvic bone ^c	13,450 ± 50	-20.7	+5.9	3.4
KIA-19643	Thoracic vertebra ^d	13,465 ± 50	-21.4	—	—
OxA-12031	Molar fragments ^e	13,205 ± 60	-21.3	+4.9	3.3
SOAN-4943	Molar fragments ^e	10,820 ± 170	—	—	—
SOAN-5063	Femur	12,970 ± 160	—	—	—
SOAN-4754	Cervical vertebrae	12,830 ± 350	—	—	—
SOAN-4753	Femur	11,840 ± 95	—	—	—
SOAN-4755	Ulna	11,310 ± 380	—	—	—
SOAN-4752	Femur	10,210 ± 135	—	—	—

^a $\delta^{13}\text{C}$ values are reported with reference to VPDB, and $\delta^{15}\text{N}$ results are reported with reference to AIR (Coplen 1994).

^bDates were run on the same sample.

^cDates were run on the same sample.

^dContains hole made by notched point.

^eUncertain if dates were run on the same specimen (see text).

Three samples of mammoth material from Lugovskoe were submitted for parallel dating to the Novosibirsk and Oxford laboratories, both to corroborate the results and to compare the effect of different collagen extraction methods. Two out of 3 samples (SOAN-4940, 4942; and OxA-12029, 12030) are in good agreement (see Table 1), but there is a serious discrepancy between SOAN-4943 and OxA-12031. When submitted, the samples were believed to be fragments of the same tooth. The samples, represented by lamellas of mammoth tooth, were collected from mud deposits. Several lamellas were used for liquid scintillation dating at Novosibirsk. The difference in the results raises the possibility that they may have come from different individuals. Alternatively, the samples may have come from the same tooth, in which case there is a significant disagreement between the 2 laboratories. Additional mammoth samples from Lugovskoe of the putative post-12,000 BP age are being submitted to resolve this important question.

Results of ^{14}C dating of the Lugovskoe mammoths (Table 1) show that bones accumulated at the site over a long period of time, from about 18,250 BP to about 10,200 BP. The 3 youngest ^{14}C values for the Lugovskoe mammoths (11,840 ± 95 BP [SOAN-4753]; 11,310 ± 380 [SOAN-4755]; and 10,210 ± 135 BP [SOAN-4752]; Table 1), are of particular importance. During the last few years, a number of mammoth dates post-12,000 BP have been reported from southern and central Siberia: Volchya Griva (about 11,090 BP), Sosva River (about 11,080 BP), and Konzul (about 11,980 BP) (Orlova et al. 2004); and from the North European Plain: Puurmani (about 10,100–10,200 BP) and Cherepovets (about 9800 BP) (Lõugas et al. 2002; Stuart et al. 2002). The contribution of the Lugovskoe date series is that it might be one of the areas with very late mammoths which survived

until the end of the Pleistocene, about 10,200 BP. In addition, a very late ^{14}C date for woolly rhinoceros from the Lugovskoe, $10,770 \pm 250$ BP (SOAN-4757) (Orlova et al., these proceedings) makes it one of the most important localities in terms of research on late survival of the Upper Pleistocene megafauna in Northern Asia.

One of the mammoth vertebrae has direct evidence of mammoth hunting by prehistoric humans, the single example known from Siberia so far. The mammoth thoracic vertebra, which is attributed to an adult female mammoth of about 23–25 yr old, has a cone-shaped hole resulting from the penetration of a notched point, and there are fragments of quartzite flakes lodged in the hole (Pavlov et al. 2002; Zenin et al. 2003). The vertebra has been ^{14}C dated to about 13,470 BP (Table 1).

CONCLUSION

The available evidence from Lugovskoe, and from several other localities in Siberia and Europe, confirm survival of woolly mammoth south of the Arctic region after about 12,000 BP, which has important implications for our understanding of the process of mammoth extinction. The Lugovskoe ^{14}C date series indicate that the nature of the mammoth habitat in Siberia after about 12,000 BP was quite “patchy,” with few isolated populations outside of the High Arctic. This is distinct from the model of “retreat to the north after approximately 12,000 BP,” which is now challenged by the new post-12,000 BP mammoth ^{14}C dates from Siberia and North Russian Plain, including the Lugovskoe. In order to test this hypothesis of extra-Arctic survival, more ^{14}C dating, including parallel dating by Russian, West European, and North American laboratories, is being undertaken. This compliments ongoing research on the chronology and paleoenvironment of mammoths and other extinct megafauna, especially in temperate Western Siberia where there may have been several megafaunal refugia after about 12,000 BP.

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