

RADIOCARBON CONCENTRATION IN THE ATMOSPHERE AND MODERN TREE RINGS IN THE KRAKÓW AREA, SOUTHERN POLAND

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ABSTRACT. New results of radiocarbon concentration in tree rings from the Kraków region covering a growth period of 20 yr have been analyzed, and the relationship between them and ¹⁴C concentrations in the atmospheric CO₂ are described. This enabled assessment of the uptake period for pine trees at the regional climatic conditions. Both sets of data show lower ¹⁴C concentrations than reported for “clean air” at the reference station, indicating a remarkable input of “dead” CO₂ of fossil fuel origin. Using data of carbon dioxide and ¹⁴C concentrations from Schauinsland, summer values of the fossil component (C_f) in carbon dioxide were calculated for the Kraków area. Fitting exponential and linear functions to experimental data, the exchange time was calculated, and expected future ¹⁴C concentration in the atmosphere was estimated.

INTRODUCTION

Recently, atmospheric concentration of carbon dioxide has exceeded 370 ppm (Keeling et al. 1995; GLOBALVIEW-CO₂ 2003), mainly due to the emission of CO₂ from fossil fuels such as coal, natural gas, and petroleum. This increase has affected global radiocarbon concentration in the atmosphere, as well as in the ocean and biosphere. Due to atmospheric nuclear bomb tests, ¹⁴C concentration in atmospheric CO₂ increased rapidly through the 1950s and 1960s to a maximum in 1963, when the measured level was double the natural value (Nydal and Lövseth 1996). Since the atomic test ban treaty was enforced, the concentration of ¹⁴C in the atmosphere has slowly decreased; however, in the 1990s, the concentration was still about 10% higher than pre-bomb levels (McNeely 1994; Levin and Kromer 1997). Several laboratories in Europe are continuously monitoring atmospheric CO₂ and/or carbon isotope composition, either directly (Levin and Kromer 1997; NOAA 2001; Necki et al. 2002; Kuc et al. 2003) or in plants (McNeely 1994; Krajcar Bronić et al. 1998; Rakowski et al. 2001). Results show that heavy local emissions of carbon dioxide from fossil fuels cause remarkable disturbances in carbon isotope composition in the local atmosphere and biosphere, and also have an influence on a regional and global scale. This effect is observed not only in heavily industrialized areas but also in urbanized regions where carbon dioxide is emitted from small industrial facilities, vehicles, and other common sources. The magnitude of the effect on ¹⁴C concentration is dependent on the distance from the sources of CO₂ emissions (Awskiuk and Pazdur 1986).

Kraków (50°3'N, 19°54'E) is a city of 1 million people, located about 100 km to the north of the Tatra Mountains in southern Poland (Figure 1). Air pollution related to the urban infrastructure, industrial facilities, and a large steel factory (Huta im. Tadeusza Sendzimiry) has elevated carbon dioxide levels in the local atmosphere (Kuc et al. 2003). Since 1983, the Department of Environmental Physics, University of Science and Technology in Kraków has monitored carbon isotope compositions in tropospheric CO₂ (Florkowski et al. 1975; Kuc 1991; Kuc and Zimnoch 1998). Samples were taken at biweekly intervals from the Faculty building roof on the university campus, 25 m above ground level. Annual tree rings were also taken from pine trees (*Pinus sylvestris*) growing near the university.

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Figure 1 Map of Poland with the location of Kraków

SAMPLES AND METHODS

In this study, 2 kinds of samples were used: CO₂ separated directly from the air at over sequential 2-week intervals and representing an integrated value over the sampling period, and tree rings obtained from trunks of pine trees (*Pinus sylvestris*) growing near the university.

Atmospheric CO₂ Samples

Samples of atmospheric CO₂ were taken by applying the method of selective sorption in capillary sorbent, as described in Florkowski et al. (1975), Kuc (1991), and Kuc and Zimnoch (1998). ¹⁴C measurements were performed with a liquid scintillation spectrometer (Tri-Carb®, Canberra-Packard) after the conversion of purified CO₂ to benzene using a standard procedure and mixing with a scintillation cocktail (b-PBD + b-MSB, Packard). Stable carbon isotope composition was measured using VG Micromass 602 C and Finningan MAT Delta S mass spectrometers at Kraków University.

Tree-Ring Samples

Annual growth rings in pine trees were sampled with a hollow drill. In order to obtain sufficient material for accelerator mass spectrometry (AMS) analysis, 3 core samples were taken from the tree and annual growth rings were separated. Samples were washed in distilled water and prepared following the standard AAA procedure. Each sample was heated at 80 °C in 0.5M HCl for 12 hr, then filtered, rinsed, and heated in 0.1M NaOH (80 °C, 12 hr). The samples were then filtered, rinsed, and heated again in 0.5M HCl (80 °C, 12 hr), then re-filtered, rinsed in distilled water, and dried. Sample residues were combined with cupric oxide, sealed in glass tubes, and evacuated by a rotary pump. The tubes were then placed in an electric furnace for 2 hr at 850 °C. Carbon dioxide produced from the samples was purified in a glass cryogenic vacuum-line system. Water produced during combustion was removed via a trap containing a mixture of methanol and liquid nitrogen at 180 °K. SO_x compounds were removed with an *n*-pentane trap cooled by liquid nitrogen. Pure CO₂ was then condensed in glass tubes at liquid nitrogen temperatures. The prepared CO₂ samples were reduced to graphite using iron powder as a catalyst (Kitagawa et al. 1993). The iron powder was prepared in a

vacuum line containing hydrogen and heated for about 1 hr at 450 °C. The line was evacuated and CO₂ was condensed with the iron powder using liquid nitrogen. Hydrogen was then added and the glass tube was sealed. Reduction was done in an electric furnace at 650 °C for 6 hr. The resulting mixture of graphite and iron powder was dried and pressed into the target holder of the AMS system. The $\Delta^{14}\text{C}$ and $\delta^{13}\text{C}$ values were measured using a HVEE Tandem AMS system and a Finningan MAT 252 mass spectrometer, respectively, at the Center for Chronological Research, Nagoya University, Japan (Nakamura et al. 2000).

¹⁴C data are reported in terms of $\Delta^{14}\text{C}$ (in ‰) deviations from the standard sample, 95% activity of NBS oxalic acid (Stuiver and Polach 1977). Stable carbon isotope data are expressed in δ -notation on the PDB scale (Craig 1957).

RESULTS AND DISCUSSION

¹⁴C concentrations in tree rings and atmospheric CO₂ from the Kraków area are shown in Figure 2, together with data from Schauinsland (Levin and Kromer 1997) as reference values. In order to compare results obtained for tree rings and atmospheric CO₂, ¹⁴C concentration data in atmospheric CO₂ were averaged for the April to September period of each year, which is the season of vegetative growth and CO₂ uptake in Poland, and corrected to $\delta^{13}\text{C} = -25\text{‰}$ to provide the $\Delta^{14}\text{C}$ notation (Stuiver and Polach 1977). Data from Schauinsland (Levin and Kromer 1997) presented in the $\Delta^{14}\text{C}$ notation were fitted to an exponential curve. ¹⁴C concentration for 1997–2002 were estimated by extrapolation on the curve. All data show a gradual decrease of ¹⁴C activity over time. Decreases of 7.1‰ per yr for tree rings and 9.15‰ per yr for atmospheric ¹⁴C were obtained. Atmospheric $\Delta^{14}\text{C}$ values measured at 2 stations in Croatia (Zagreb and Plitovice) during the same period decrease at rates of 12‰ and 10.6‰ per yr, respectively (Krajcar Bronić et al. 1998). For the Schauinsland station, decreases of 14.1‰ per yr for the 1983–1985 period and 9.7‰ per yr for 1985–1989 were reported (Levin et al. 1992).

Absolute differences between $\Delta^{14}\text{C}$ values from tree rings and atmospheric CO₂ from the Kraków area are presented in bar form in Figure 2. ¹⁴C concentrations in atmospheric CO₂ show strong peak-to-peak (~56‰, Kuc and Zimnoch 1998) and seasonal fluctuations that are not recorded in the tree-ring samples. ¹⁴C concentrations in both kinds of samples from Kraków are lower than in “clean air” at the Schauinsland station. This is attributed to a local Suess effect. The $\Delta^{14}\text{C}$ values in tree rings (108‰, 99‰, and 85‰ for 1994, 1995, and 1996, respectively) are lower than the mean yearly values in atmospheric CO₂ at Schauinsland (117‰, 111‰, and 102‰, respectively; Levin and Kromer 1997) but higher than those measured in Zagreb (81‰, 88‰, and 56‰; Krajcar Bronić et al. 1998) for the same period.

Exponential curves were fitted to the Kraków data for the period of 1983–2002 to determine long-term changes in ¹⁴C concentration in the atmosphere and biosphere. These analyses provide time decay constants of 16.7 yr for tree-ring data and 14.3 yr for atmospheric CO₂, which correspond well with ~16-yr estimates reported in the literature (Levin et al. 1995; McNeely 1994; Krajcar Bronić et al. 1998). Carbon isotope ratios calculated from exponential and linear models decrease at rates of 7.1‰ and 6.75‰ per yr, respectively, for tree-ring data, and 7.75‰ and 6.29‰ per yr for atmospheric CO₂ data. According to the linear model, $\Delta^{14}\text{C}$ is estimated to equal 0‰ in 2010 for tree rings and 2007 for atmospheric CO₂. The same procedure produces estimates of 2007 for the Schauinsland data (Levin and Kromer 1997) and 2004 for the Groningen data (Meijer et al. 1995).

¹⁴C concentration in the atmosphere over large cities is strongly affected by emissions of CO₂ from fossil fuels. Concentrations of carbon dioxide can be divided into 3 components: a background com-

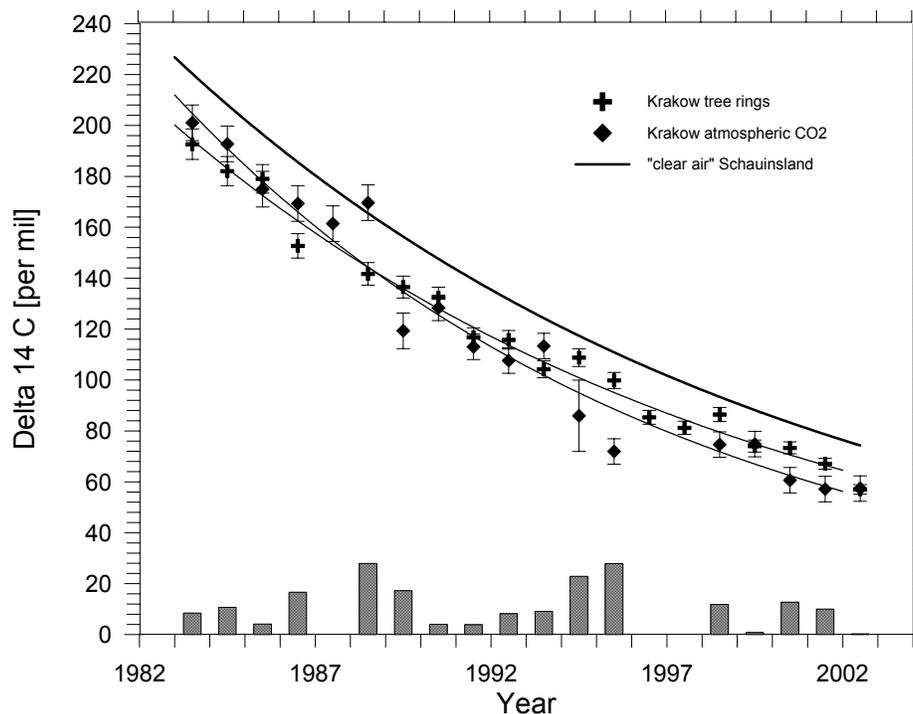


Figure 2 $\Delta^{14}\text{C}$ values in tree rings and atmospheric CO_2 samples collected in the Kraków area, southern Poland. The solid line represents ^{14}C concentration in “clear air” at Schauinsland station, estimated from the measurement of atmospheric CO_2 (Levin and Kromer 1997). Bars show absolute differences between atmospheric CO_2 and tree-ring samples for Kraków. All atmospheric data represent mean annual values for the April to September season, expressed as $\Delta^{14}\text{C}$.

ponent (C_a), a biogenic component (C_b), and a fossil component (C_f). Mathematical equations that can describe the relationships between each component and carbon isotopic composition are discussed in previous literature (e.g. Levin et al. 1989; Zondervan and Meijer 1996; Kuc et al. 2003), and were used to calculate C_f for the Kraków data. The value of C_f varies seasonally, from ~ 27.5 ppmv in the winter to ~ 10 ppmv in the summer during the 1983–1994 period (Kuc and Zimnoch 1998; Kuc et al. 2003). This variation correlates with differences in fossil fuel consumption between the seasons in the Kraków area.

New estimates of C_f using the new Kraków data are presented in Figure 3. Reference values of ^{14}C and CO_2 concentration representing “clean air” were calculated using data from Schauinsland station (GLOBALVIEW- CO_2 2003). Averaging C_f over 2 decades (1983–2003) for the April to September season, values of 5.9 ppmv and 6.5 ppmv were obtained for tree rings and atmospheric CO_2 , respectively. Both values are lower than the yearly average obtained for consecutive years by Kuc and Zimnoch (1998); however, the estimates converge over time. Lower values of C_f observed in the Kraków area in recent years are the result of a reduction in fossil fuel consumption, especially in the case of coal.

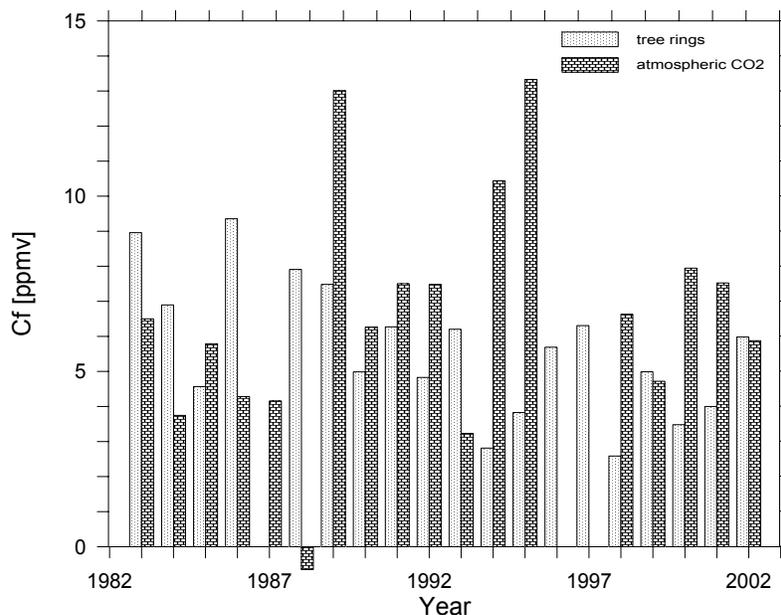


Figure 3 Values of the fossil component C_f calculated using tree-ring and atmospheric CO_2 data from the Kraków sampling site. Data represent mean annual values for the April to September season.

CONCLUSION

¹⁴C concentration in samples of tree rings and in atmospheric CO_2 , measured in Kraków over last 20 yr, are lower than in “clean air” due to the input of CO_2 from fossil fuels. Decreases in ¹⁴C concentration of 7.1% per yr for tree rings and 9.15% per yr for atmospheric CO_2 were observed. Data fitted to exponential and linear functions provided reliable estimates of rates of decrease in carbon isotope ratios and time decay constants for the sampling area. Isotopic records in tree rings and atmospheric CO_2 collected over long time periods yield information on the local input of anthropogenic CO_2 . Moreover, tree rings appear to be a valuable archive of isotopic variations in the local environment over time, albeit in annual increments only. In comparing atmospheric CO_2 and tree-ring data, it is necessary to consider the climatic conditions of central Europe, since the commencement and duration of the vegetative growth (and CO_2 uptake) period can vary from year to year.

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REFERENCES

Awskiuk R, Pazdur MF. 1986. Regional Suess effect in Upper Silesia urban area. *Radiocarbon* 28(2A):655–60.
Craig H. 1957. Isotope standards for carbon and oxygen and correction factors for mass-spectrometric analysis of carbon dioxide. *Geochimica et Cosmochimica Acta* 12:133–49.
Florkowski T, Grabczak J, Kuc T, Rózański K. 1975. Determination of radiocarbon in water by gas or liquid

- scintillation counting. *Nukleonika* 20(11–12):1053–66.
- GLOBALVIEW-CO₂. 2003. Cooperative Atmospheric Data Integration Project—Carbon Dioxide. CD-ROM, NOAA CMDL, Boulder, Colorado [Also available on internet via anonymous FTP to ftp.cmdl.noaa.gov, Path: ccg/co2/GLOBALVIEW].
- Keeling CD, Whorf TP, van der Plicht J. 1995. International extremes in the rate of rise of atmospheric carbon dioxide since 1980. *Nature* 375:666–9
- Kitagawa H, Masuzawa T, Nakamura T, Matsumoto E. 1993. A batch preparation method for graphite targets with low level background for AMS ¹⁴C measurements. *Radiocarbon* 35(2):295–300.
- Krajcar Bronić I, Horvatinčić N, Obelić B. 1998. Two decades of environmental isotope records in Croatia, reconstruction of the past and prediction of future level. *Radiocarbon* 40(1):399–416.
- Kuc T. 1991. Concentration and carbon isotope composition of atmospheric CO₂ in southern Poland. *Tellus* 43B:373–8.
- Kuc T, Zimnoch M. 1998. Changes of the CO₂ sources and sink in polluted urban area (southern Poland) over last decade, deriving from the carbon isotope composition. *Radiocarbon* 40(1):417–23.
- Kuc T, Rozanski K, Zimnoch M, Necki JM, Korus A. 2003. Anthropogenic emissions of CO₂ and CH₄ in an urban environment. *Applied Energy* 75:193–203.
- Levin I, Schuchard J, Kromer B, Münnich O. 1989. The continental European Suess effect. *Radiocarbon* 31(3):431–40.
- Levin I, Kromer B. 1997. Twenty years of high-precision atmospheric ¹⁴CO₂ observation at Schauinsland station, Germany. *Radiocarbon* 39(2):205–18.
- Levin I, Graul R, Trivett NBA. 1995. Long-term observations of atmospheric CO₂ and carbon isotopes at continental sites in Germany. *Tellus* 47B:23–34.
- Levin I, Böisinger R, Bonani G, Francey RJ, Kromer B, Münnich KO, Suter M, Trivett NBA, Wölfli W. 1992. Radiocarbon in atmospheric carbon dioxide and methane: global distribution and trends. In: Taylor RE, Long A, Kra RS, editors. *Radiocarbon After Four Decades: An Interdisciplinary Perspective*. New York: Springer-Verlag. p 503–18.
- McNeely R. 1994. Long-term environmental monitoring of ¹⁴C levels in Ottawa region. *Environment International* 20(5):675–9.
- Meijer HAJ, van der Plicht H, Gislöfoss JS, Nydal R. 1995. Comparing long-term atmospheric ¹⁴C and ³H records near Groningen, the Netherlands with Fruholmen, Norway and Izaña, Canary Islands ¹⁴C stations. *Radiocarbon* 37(1):39–50.
- Nakamura T, Niu E, Oda H, Ikeda A, Minami M, Takahashi H, Adachi M, Pals L, Gott dang A, Suya N. 2000. The HVEE Tandatron AMS system at Nagoya University. *Nuclear Instruments and Methods in Physics Research B* 172:52–7.
- Necki J, Schmidt M, Rozanski K, Zimnoch M, Korus A, Lasa J, Graul R, Levin I. 2002. Six-year record of atmospheric carbon dioxide and methane at a high-altitude mountain site in Poland. *Tellus* 55B:94–104.
- NOAA Climate Monitoring and Diagnostics Laboratory. 2001. ftp site: ftp://cmdl.noaa.gov/ccg/network.txt.
- Nydal R, Lövsæth K. 1996. Carbon-14 measurement in atmospheric CO₂ from Northern and Southern Hemisphere sites, 1962–1993. Oak Ridge National Laboratory NDP-057.
- Rakowski AZ, Pawelczyk S, Pazdur A. 2001. Changes of ¹⁴C concentration in modern trees from Upper Silesia region, Poland. *Radiocarbon* 43(2B):679–89.
- Stuiver M, Polach HA. 1977. Discussion: reporting of ¹⁴C data. *Radiocarbon* 19(2):355–63
- Zondervan A, Meijer AJ. 1996. Isotopic characterization of CO₂ sources during regional pollution events using isotopic and radiocarbon analysis. *Tellus* 48B:601–12.