A STUDY OF ATMOSPHERIC DEPOSITION ONTO THE SNOWPACK IN NORTHERN SASKATCHEWAN

by

S..R. Shewchuk

(Saskatchewan Research Council, 30 Campus Drive, Saskatoon, Saskatchewan S7N 0X1, Canada)

ABSTRACT

The purpose of this study is to use the snowpack the Saskatchewan portion of the Precambrian Shield of as a passive collector of atmospheric total deposition. Although the focus is on the anions of sulphate and nitrate, the hydrogen ion and various metals are also considered.

Twenty-six small lakes on the Precambrian Shield area of northern Saskatchewan were used as sites. The lakes were all approximately 1 square km in area and serviced by aircraft. The study was conducted in late March 1982 and 1983 just before the major snowmelt period.

Concentrations of sulphate ranged from 0.26 to 7.3 mg l⁻¹. The concentrations were typical of large areas in northern Saskatchewan with sharp transitions in composition occurring as the treeline boundary was encountered. Typical sulphate deposition was approximately 30 mg m^{-2} in the Precambrian areas and 150 mg m^{-2} in the other areas. The distribution of conductivity in the snow cores ranged from 6 to 13 µS cm⁻¹. Most of the snow on the Precambrian Shield areas can be described as acidic, with pH values of the meltwater being below 5.6.

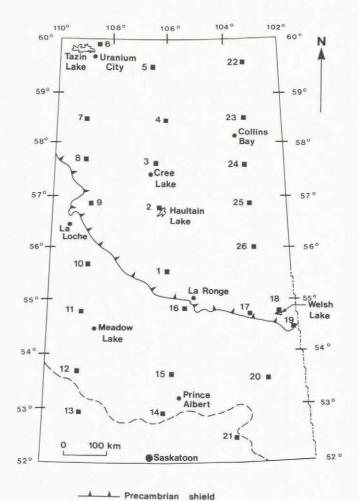
Metals such as Ca, Mn, Mg, Zn and As were measured across the region. The metals Ca and Mg show a direct correlation with high sulphate levels south of the treeline. Metals such as Zn and As show anthropogenic influence particularly within 100 km of industrial sources. Large areas of the study region show uniform and low deposition to the snowpack. It is suggested that natural atmospheric processes are dominating deposition in this area of Canada.

INTRODUCTION

The purpose of this study was to use the snowpack of northern Saskatchewan as a passive collector to determine the extent of anthropogenic deposition to the various ecosystems. Over northern Saskatchewan a large portion of the yearly precipitation falls as snow. Provided air temperatures are below freezing, snow surfaces trap gases and particulates (Elgmork and others 1973, Barrie 1980). During a winter with little or no snowmelt most of the pollutant load is retained in the snowpack. According to Johannessen and Henriksen (1978), from 50 to 80% of the pollutant is released in the first 30% of meltwater.

Shewchuk (1979) has shown that trace levels of sulphur dioxide are entering the Precambrian region of Saskatchewan on an episodic basis. It is quite likely that this sulphur dioxide is influencing acid deposition in this area.

Snow surveys were conducted on a series of small lakes in northern Saskatchewan during late March in the years 1982 and 1983. The study area, within western Canada, is shown in Figure 1. Twenty-six small lakes were selected to cover this study area uniformly, representing approximately $4.6 \times 10^5 \text{ km}^2$. The area is composed of granitic and quartzite sandstone bedrock and is largely devoid of carbonate. On the Precambrian Shield, there are many freshwater lakes; these have been



---- Treeline boundary

Fig.1. Snow survey site locations within the study area.

classified as being highly sensitive to acid deposition (Shewchuk 1982).

There are many large point source emitters within 100 km of the boundary of the study area. At present they account for in excess of 346000 tonnes a^{-1} of sulphur dioxide emission (Canada. Minister of Supply and Services 1981). Part of this emission originates from a metal smelter at Flin Flon, Manitoba (212000 tonnes a-1) while the oil sands industry located near Fort McMurray, Alberta contributes much of the rest (134000 i). Significant increases in output of emissions tonnes a of sulphur oxides is expected in the next several decades (Webber and Warne 1979). The oil sands area is located predominantly upwind of the highly sensitive Precambrian Shield (Hopkinson and Dublin 1981) of Saskatchewan.

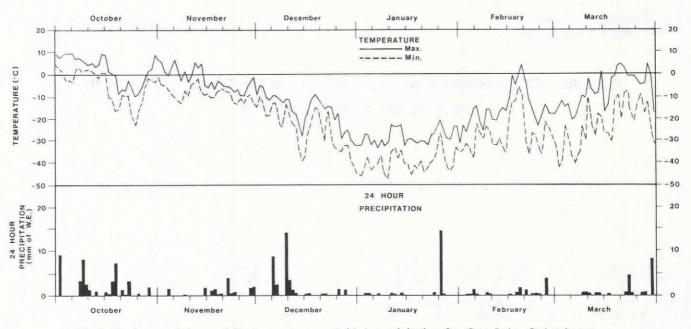


Fig.2. Maximum-minimum daily temperatures and 24 h precipitation for Cree Lake, Saskatchewan, October 1981 to March 1982. (Data from Environment Canada.)

PROCEDURES

The twenty-six lake sites were reached by a ski-equipped Cessna 185 aircraft. Lakes selected for this study had to be approximately 1 km^2 in area and be representative of the region in which they were located. The small lake allowed for convenient access and minimized the effect of snow drifting off the lake surface. Nearly all these sites were remote and only accessible by aircraft.

Normally the centre of the lake surface proved to be the best location from which snow cores could be taken. Snow cores were taken with acrylic tubes 1.5 m in length with diameters of 10 and 15 cm, two sizes being used so that the snow volume could be controlled (i.e., for deep snow the smaller diameter would be used and for shallow snow the larger diameter would be used). Normally the corer was inserted vertically into a clean, smooth area of snow. It was pushed through the snowpack to the surface of the lake ice. It was then dug out with a plastic shovel. The texture of the ice surface was noted carefully and if any slush or melting was detected the snow core was terminated just above the slushy area. It was important to avoid the collection of surface lake slush and water because it is the chemistry of the snowpack that is the subject of this study and not the chemistry of a snow-water interface. The presence of meltwater on the ice surface below the snowpack was rare. The ice surface at the base of the snowpack was unusually hard, smooth ice. The chemical composition of the total core was of interest. However, if major melt periods identified as layering in the snow observed the separate layers were collected were individually.

The samples were transported to the laboratory in the frozen state. They were kept frozen up until the time of chemical analysis. The snow was allowed to melt and as soon as room temperature was reached the sample was analyzed. Generally sample analysis was completed on the same day. Sulphate was determined in the melted sample by ion chromatography. Nitrate was determined by cadmium reduction, and pH was determined both by electrode techniques and Gran (1952) titration procedures. With the exception of arsenic, metals were determined by inductively coupled plasma emission spectroscopy. Arsenic was determined by hydride generation.

For most lake surfaces three snow cores were analyzed for all constituents. At one location detailed deposition studies were carried out on land snow cores as well as lake snow cores in an attempt to assess the effect of drifting.

RESULTS

Winter in this area is an ideal time for the determination of the extent of anthropogenic deposition on to the snowpack by means of snow chemistry studies. The maximum-minimum temperature structure and the 24 h precipitation at Cree Lake is shown in Figure 2. Cree Lake is located centrally in the region. The air temperature is consistently below 0°C from about mid-November, and little melting occurs until mid-March. It is therefore reasonable to assume that most of the chemical deposition arriving on the snow surface remains essentially within the snowpack during the period from mid-November to mid-March.

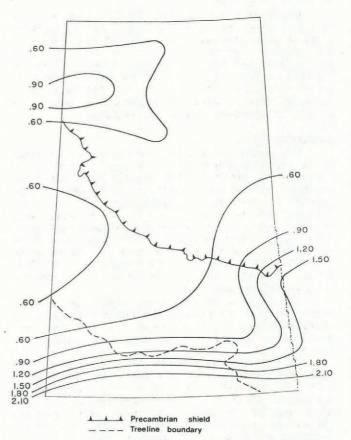


Fig.3. Sulphate concentration (mg 1⁻¹) in snowmelt water collected in March 1982.

6		Tazin Lake	Haultain Lake	Welsh Lake
Ca	(mg l ⁻¹)	0.16	<0.05	1.0
C1	(mg l ⁻¹)	0.03	0.01	0.066
HCO3	(mg l ⁻¹)	0.6	<0.5	3.9
К	(mg 1 ⁻¹)	<0.1	<0.1	0.2
Mg	(mg 1 ⁻¹)	0.01	0.02	0.48
Na	(mg 1 ⁻¹)	<0.1	<0.1	0.2
SO4	(mg l ⁻¹)	0.67	0.46	1.7
N(NH ₃)	(mg l ⁻¹)	0.04	0.04	0.04
N(NO ₃)	(mg l ⁻¹)	0.12	0.16	0.18
Specific conductivity (μ S cm ⁻¹)		4.8	7.4	12.0
Total volume (ml)		725	500	990
pH (electr	ode)	5.12	4.81	7.40
pH (Gran's plot)		5.02	4.76	nil

TABLE I. CONCENTRATION IN THE SNOW OF VARIOUS IONS AT TAZIN LAKE, HAULTAIN LAKE, AND WELSH LAKE

TABLE II. SNOWMELT WATER CHEMISTRY AND SULPHATE DEPOSITION ONTO THE LAKE AND LAND SURFACE OF THE SITE AT HAULTAIN LAKE FOR THE MARCH 1983 STUDY

Surface	Snow depth (cm)	Meltwater (ml)	Equivalent water depth (cm)	Conductivity (µS cm ⁻¹)	pH (range)	Sulphate concentration (mg 1 ⁻¹)	Sulphate deposition (mg m ⁻²)
Lake	28±4	529±62	6.5±0.7	6.9±0.6	4.81-6.75	0.47±0.06	30.6
Land	46+10	858±132	10.6±1.6	6.7±0.4	4.43-4.93	0.47±0.09	49.8

TABLE III. COMPARISON OF SNOWMELT WATER CHEMISTRY PARAMETERS FOR HAULTAIN LAKE BETWEEN 1982 AND 1983

Year	SO ₄ ²⁻ (mg l ⁻¹)	NO ₃ -N (mg l ⁻¹)	Conductivity (µS cm ⁻¹)	Snow density	pH (range)
1983	0.47±0.06	0.15±0.01	6.9±0.6	0.23±0.04	4.81-6.75
1982	0.43±0.07	0.09±0.01	5.7±1.8	0.24±0.03	5.04-6.78

Distance SO4 Fe Mn Zn Pb As (km)5 1700 110 2 36.0 720 84 15 1400 400 5 23.0 400 35 9 30 750 60 2 2.6 70 970 170 8 2.0 29 4 60 890 22 2 1.1 13 5 100 300 470 34 10 < 0.5 6 4 490 1 < 0.5 5 <2 600 24

TABLE IV. CONCENTRATIONS OF TRACE ELEMENTS (in 10⁻⁹ g g⁻¹) IN THE SNOW CORES AT VARIOUS DISTANCES FROM THE SMELTER AT FLIN FLON ALONG A TRANSECT TO THE NORTH-WEST

The regional distribution of sulphate concentration in the melted snow cores is shown in Figure 3. In the vast area of the Precambrian Shield the sulphate distribution is reasonably uniform. The sharp transition in concentration occurs when the area south of the treeline is encountered. This area is mostly farm land. Snow drifts and mixes with the soil dust quite extensively. Consequently most snow cores consist of well-mixed constituents of soil dust. In the forested areas north of the treeline, however, the soils are much less exposed to wind erosion. The intrusion of sulphate into the south-east corner is likely from the metal smelters at Flin Flon, in combination with natural influences. The natural contribution is clearly defined because increases in sulphate in this region of the study area are paralleled by increases in alkaline earth metal concentrations (Table I). Welsh Lake which is located in the higher sulphate zone of the south-east also has much higher levels of calcium and magnesium than does the Haultain Lake area. The intrusion into the north-west corner is probably a natural impact of soil dust from the Athabasca River valley. The Athabasca River valley is approximately 50 km adjacent to the Precambrian Shield area in this part of Canada and is least sensitive to acid deposition.

The concentration of various constituents across the Precambrian Shield is shown in Table I. The lakes represented by the Tazin and Haultain site are considered typical of central shield areas. While Welsh Lake is located on the shield, its more southerly location shows higher deposition for every constituent with the exception of hydrogen ions. Acidity values in the central shield range from 20 to 40 μ eq⁻¹. The amount of acidity decreases rapidly at Welsh Lake. This site is located 50 km west of the smelter source at Flin Flon. It should be noted that the concentration of nitrogen compounds such as ammonia and nitrate are exceedingly low in this study area.

The bicarbonate concentration at lakes located in the central portion of the shield is near detection limits. However as the southern fringe areas of the shield areas are entered, the bicarbonates increase in concentration. Elements such as Ca, Mg, K and Na are also seen to increase in concentration as the southerly area is approached. The pH values determined by electrode and titration procedures were in good agreement provided there was sufficient hydrogen ion present (>10 μ eq 1⁻¹). However as the hydrogen ion content was lowered the titration methods were of limited value, for example, at the Welsh Lake site.

Table II is a summary of observations of land and lake depositions at the Haultain Lake site which is located near the centre of Saskatchewan's Precambrian Shield. Each observation represents the mean and standard error of 16 separate observations taken from each surface. The lake basin is typical of the freshwater ecosystems that are highly sensitive to acid deposition. However the concentration of sulphate in the meltwater remains similar.

A comparison of concentrations of sulphate, nitrate, hydrogen ion, etc., at Haultain Lake is shown in Table III for the two winter seasons where data are available. While increases in concentration of the order of 10% are observed, the total sulphate and nitrate concentrations are not significantly different within the limits of variability.

The variation in concentrations of sulphate and metals in snow with distance from the Flin Flon smelter are shown in Table IV. Most metal concentrations are quite low on the shield, often approaching values that are typical of remote regions. The concentration of many of the anthropogenic metals, such as As, Zn, and Pb, is influenced by the smelter for a distance of at least 100 km. The gradient of arsenic is particularly evident. It is seen that elements that are more crustal in origin (Fe and Mn) have a more uniform distribution along the transect.

CONCLUSIONS

The regional distribution of sulphate in the meltwater of the snow from northern Saskatchewan is low and relatively uniform. The influence of open prairie land on the concentration of sulphate in the snow suggests that the soil is contributing substantially to this ion in the snowpack. However, in areas covered by trees remote from prairie lands deposition is likely to come directly from the atmosphere. The total deposition of sulphate to the snow of the shield during winter is well below 100 mg m⁻², and the pH of the snow is ~4.9. Recent evidence suggests that organic acids may influence to some extent the acidity of precipitation in remote areas (Keene and others 1983).

ACKNOWLEDGEMENTS

Major funding for this research was provided by Saskatchewan Environment. I would like to thank Mr L J Lechner of Saskatchewan Environment for his interest in my group. In addition, special acknowledgements are due to several employees of the Saskatchewan Research Council. They are Mr Bob MacPherson, who was the aircraft pilot, and Mr James Ross, who was the laboratory assistant in this project. REFERENCES

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